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THE GEOMAGNETIC FIELD
OF THE NETHERLANDS
REDUCED TO
1945.0

BY

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PREFACE

ON publishing this report I wish to express my gratitude to all persons who have taken part in the magnetic survey of the Netherlands. In the first place thanks are due to Mr. J. OLDEMAN, scientific assistant at the 5th department of the Royal Netherlands Meteorological Institute for his assistance and cooperation, to the staff of the magnetic observatory Witteveen, and to the other members of the staff of the 5th department, who undertook the troublesome work of reducing the measurements at De Bilt as well as at Witteveen, and further to Dr. C. P. HARTMANN, seismologist at the Bataafse Petroleum Maatschappij, who performed the greater part of the Z-measurements during the years 1942 to 1944. Moreover thanks are due to Messrs. Dr. P. GROEN, A. HAUER, Dr. H. TEN KATE, Dr. W. F. SCHALKWIJK and C. G. C. SCHÜTTE, members of the scientific staff of the Royal Netherlands Meteorological Institute, who bestowed part of their time on the survey during wartime. I wish to acknowledge the great amount of accurate work done by them. Finally I am greatly indebted to the commander of the Netherlands Navy, Admiral J. J. L. WILLINGE, who several times put a ship at my disposal for the measurements in the Wadden Sea.

De Bilt, July 1950

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INTRODUCTION

THIS report contains the results of geomagnetic measurements carried out in the summers of the years 1942—1948. The idea of measuring the geomagnetic field of the Netherlands was suggested to the present author at the end of 1941 by Mr. C. P. HARTMANN, at that time cand. geol. From the first discussions a broader plan arose, which aimed at a detailed survey to be carried out by the Royal Netherlands Meteorological Institute. A number of members of the scientific staff, who could not perform their normal duties owing to war circumstances, were able to give their cooperation in making the measurements. They were Messrs. GROEN, HAUER, TEN KATE, SCHALKWIJK and SCHÜTTE. The assistance of Mr. OLDEMAN of the geophysical department must be mentioned especially. Mr. HARTMANN took part as a guest. The whole survey was under the direction of the present author.

It will be evident, that the war hampered this work. In different parts of the Netherlands, especially in the coastal region, fieldwork was forbidden by the Germans. The problems connected with transport, food and lodging of the measuring party were considerable, and in the last year of the war measurements were of course impossible.

During the summer months of the years 1942, 1943 and 1944 measurements were carried out in a great part of the Netherlands, with the exception of the coastal regions, the Zuid-Hollandse islands, Zeeland, Zeeuws-Vlaanderen and the Wadden islands. The results of these measurements have been used by HARTMANN for the composition of his thesis¹⁾.

After the liberation in 1945 the survey could be resumed, thanks to the fact, that the instruments and data had all been saved and that, moreover, the magnetic observatory at Witteveen had sustained the liberation without any damage. The measurements in the years 1945 till 1948 were carried out almost exclusively by Mr. OLDEMAN and the author. The country of Zeeland, Zeeuws-Vlaanderen and the Zuid-Hollandse islands was measured in the fall of 1945 and the summer of 1946. In the summer of 1947 measurements were made on the Wadden islands and on some sand-banks in the Wadden Sea. The Commander of the Netherlands Navy, the Admiral J. J. L. WILLINGE put a small yacht, Hr. Ms. Neptunus (RC 39) with crew under the kwartiermeester RUIJSAARD at our disposal during three weeks for transporting the instruments and the observers. When it became clear that an important anomaly was present in the Wadden Sea, the desirability was felt, to localize this anomaly more exactly by a number of supplementary observations. These were carried out in the summer of 1948. Support was asked from the Commander of the Netherlands Navy, who kindly lent the naval craft Doornbosch (RQ 2) for three weeks. Captain of the craft was Schipper VAN DER MEY. With the aid of this vessel fourteen measurements were performed on sand-banks in the Wadden Sea. The outcome of these observations has been given in a separate table.

¹⁾ PH. C. P. HARTMANN, Aardmagnetische Anomalieën in Nederland, Utrecht 1945.

They have been used for the construction of the magnetic charts of the components H , Z and ΔZ .

The geomagnetic field of the Netherlands had already been investigated previously. In the years 1889 to 1892 measurements were carried out by E. VAN RIJCKEVORSEL¹⁾ at more than 300 places in our country. The result was laid down in a bulky report: A magnetic survey of the Netherlands, for the Epoch January 1, 1891, edited in the series "Nieuwe Verhandelingen van het Bataafs Genootschap der Proefondervindelijke Wijsbegeerte te Rotterdam". Three components of the geomagnetic field were investigated by him: the declination, the horizontal force and the inclination. For the first and the second VAN RIJCKEVORSEL used an ELLIOT unifilar magnetometer, for the determination of the dip a DOVER dip-circle. All the measurements were performed by VAN RIJCKEVORSEL alone, a few times he was aided by a servant. Travelling was mostly done by train, less often by boat or carriage. Especially the travelling by train has strongly influenced the choice of many measuring-places. He writes about this matter:

"I do not at all take much care in selecting my station. As soon as I arrive at a place where I wish to observe, I put down my instruments on the first spot which looks likely to be good, without making a severe hunt for iron, or taking many informations about gaspipes or rails. Even, as many railway-stations, especially of the state-railways, have a large tract of enclosed ground, I very often observed there. I thus had the advantage of finding, without loss of time, a convenient spot, where there were no troublesome intruders, but of course I often came dangerously near to rails or waggons. I was quite aware of this objection to my method. But it gave me the great advantage of gaining time. It will be seen that I regularly took two stations a day and in one or two cases even more than that."

Owing to this method he worked close to the rails or other large iron masses in many cases. This becomes evident from the remarks given in the description of the stations, like: "20 m of corner of station", "near a shed containing a lot of machinery", "during part of work had to suffer much from smoke of locomotive". Moreover, VAN RIJCKEVORSEL used to choose his standpoint near a building. Repeatedly we find descriptions like: "garden of inn", "garden of stationmaster", "garden of hotel", "in front of the only house near the station", "market-place", "within the town", "near front of old lighttower", "near church, within enclosure".

Very exclusive indeed was the measurement carried out at Nesserzand, a sand-bank easterly of the island of Texel: "In the sea with water nearly up to the waist. Geographical position kindly communicated to me by the commander of the torpedoboat, which had been lent to me. For the sun's observation, as I could not have the chronometer near me, the commander had it in the boat and gave the time for each single reading".

¹⁾ Dr E. VAN RIJCKEVORSEL, A magnetic survey of the Netherlands for the epoch January 1, 1891 (Rotterdam 1895).

Out of the 328 measurements about 130 have been performed at places, which decidedly must be called bad from a magnetic viewpoint, whereas about 50 more are less favorable. Only 150 places can be qualified as good, judging from the description.

As a consequence of this lack of care in choosing the place where the instruments were mounted, a great number of the measurements (about a half) have been carried out in places which are magnetically disturbed, and therefore give an untrue picture of the magnitude and direction of the magnetic field. Besides, the present author cannot throw off the impression that the accuracy of the measurements is smaller than given by VAN RIJCKEVORSEL. In his discussion about the declination measurements for example, he estimates the accuracy of determining the magnetic meridian at $\pm 3''$. Our impression is that no accuracy better than $10''$ can be obtained with the instrument used by VAN RIJCKEVORSEL. The exactness of the chronometer, which is determinative for the accuracy of the geographic north, is estimated at ± 2 sec, although checking the rate of the chronometer took place only at the beginning and at the end of a 14 days' travel. Our experience with similar chronometers, which were compared with wireless time-signals every day, are much worse however. Finally the determinations of the vertical force, calculated from the horizontal force and the inclination, have been affected by the use of the dip-circle. This is an inaccurate instrument, as the position of the needles is apt to be changed by the slightest irregularities of the axis of rotation. The possible error of the inclination is estimated by VAN RIJCKEVORSEL at about $1'$, which corresponds to an inaccuracy of about 35% in the vertical force. It is quite certain that in many measurements the error has been larger than this figure.

The result of all is, that the geomagnetic field in the Netherlands has a highly disturbed character after VAN RIJCKEVORSEL's measurements. The differences between values at adjacent stations are most irregular. As a result the isomagnetic lines are drawn in capricious windings on his charts. This is especially the case with the chart of the vertical force. As this chart is of great importance for a geological interpretation of the irregularities, it seemed worth while to repeat the observations with more accurate methods. Besides, the geomagnetic field has considerably altered since 1891 by the secular change, so that a new survey was desirable.

CHAPTER I

INSTRUMENTS AND METHODS OF MEASUREMENT

The members of the staff who took part in the survey, joined the magnetic party alternately. In the years 1942 and 1943 the party was composed mostly of three persons. Later on it turned out, that the work could be done by two persons quite as well.

If the number of participants was three, each of them was occupied with one component of the geomagnetic field. Consequently there were participants who were specialized in measuring the declination D , others who were accustomed to the determination of the horizontal component H , again others for the vertical component Z . This distribution

of work proved to be an advantage for the quality of the measurements.

When the party consisted only of two persons, that distribution of work was the best, in which one of them carried out one D -determination together with one Z -measurement, whereas the second combined the H -measurement with two more Z -determinations.

Transport of the party and the instruments took place by motorcar. The carrying out of one complete measurement D , H and Z took about two hours, and when the weather was unfavourable so that wind-screens had to be put up, three hours or more were necessary. Under favourable conditions it was possible to do three stations a day. At each station one D -measurement was carried out, H was determined at two places and Z at three places, the distances between two places being 20 to 50 m. This method had the advantage that already during the measurements an impression was got about the magnetic properties of the ground. In cases where the Z -observations differed strongly, it was supposed that a local disturbance was present and the station was moved to a better place. Such disturbances were discovered several times. In general much care was taken to avoid artificial disturbances. Minimal distances were fixed for diverse iron or iron containing objects. These minimal distances were: instrument case 5 m, motorcar 25 m, barbed wire 5 m, iron fence 30 m, iron pipe line 25 m, rails 200 m, mine shaft 500 m, buildings 100 to 500 m, ballasted roads 100 m. These distances were partly taken from "Physics of the Earth VIII", page 132, partly they were determined by experiment. Preferably the instruments

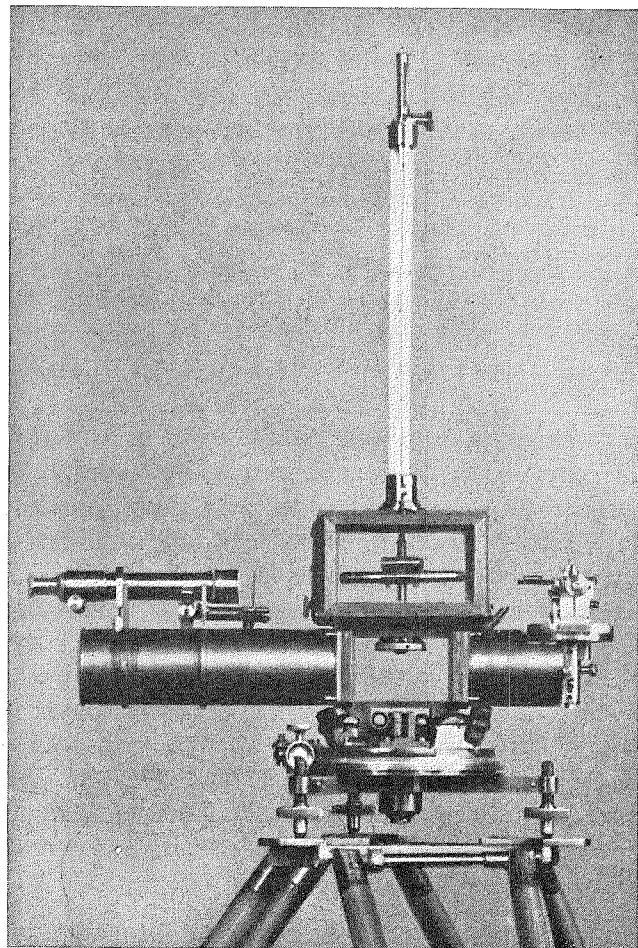


Fig. 1. Theodolite for declination measurements.

were set up in meadows, on arable ground or uncultivated sandy ground. In order to avoid mutual disturbances minimal distances of 15 m were kept between the instruments.

All measurements of the magnetic elements were made at the moment of the minute-record in the magnetic observatory Witteveen, where every minute, a one-second snapshot of the geomagnetic field is taken.

Measuring the declination.

The instrument (fig. 1) was a theodolite after ELLIOT left by VAN RIJCKEVORSEL, which permitted us to read directions with an accuracy of 10". With a view to the declination-measurements several changes were introduced. To enable the triangulations to be carried out, a little telescope (with an object-diameter of 25 mm) was fixed. For the astronomical measurements the mirror as well as the mirror's axis were constructed with the utmost accuracy.

First a description will be given of a declination-measurement by means of a sun's observation.

The theodolite is mounted on a tripod as firmly as possible. Then the axis of rotation is adjusted vertically by levelling. The axis of the mirror must then be horizontal, and this can be checked by a levelling instrument. If necessary the position of the axis is corrected. Besides, the mirror's axis has to be perpendicular to the optical axis of the telescope, and the reflecting surface of the mirror has to be parallel to the mirror's axis. These conditions cannot be checked in the field, for this requires a dark room. The checking was carried out several times a year in the observatory. The reticles of the telescope are illuminated from one side and the mirror is manipulated in such a way that the reflection of the reticles coincides with the object. It goes without saying that the vertical reticle must be vertical indeed, when the angles between objects of different altitudes like the image of the sun and church-steeple are measured.

Now an image of the sun is projected in the telescope by help of the mirror. Two small blue glasses can be placed in the ray path in order partially to absorb the sunlight. Then several times the moment is read, at which the left and right limbs of the sun touch the reticle-wire. This can be done in the most simple way by an assistant counting the seconds of the chronometer. At the same time the azimuth of the theodolite is read. This series of readings is repeated after the mirror has been lifted out of the supports, turned halfway round its vertical axis and let down in the supports again. This enables us to eliminate the error introduced by the mirror's axis and the mirror not being parallel. An eventual error in the position of the mirror's axis with respect to the axis of the telescope is removed by a determination with the sun in front of the observer as well as behind him. If it turns out that in the first case the direction of the astronomical north is read by an angle A , and in the second case by an angle B , the true azimuth N is given by¹⁾

$$N = \frac{A + B}{2} + \frac{A - B}{2} \cos h,$$

¹⁾ PH. C. P. HARTMANN, Thesis Utrecht 1945.

where h is the elevation of the sun, computed from the formula

$$\sin h = \sin \varphi \sin \delta + \cos \varphi \cos \delta \cos \tau$$

or from

$$\cos h = \frac{\cos \delta}{\sin \alpha} \sin \tau.$$

In these formulas φ is the geographical latitude of the station, δ is the declination of the sun, τ is the local true hour-angle of the sun, and α is the sun's azimuth. The azimuth of the sun is computed from

$$\operatorname{tg} \alpha = \frac{\sin \tau}{\sin \varphi \cos \tau - \cos \varphi \operatorname{tg} \delta}$$

The hour-angle is derived from the true local time, calculated from the observed Greenwich-time, plus the longitude of the station converted in time, and minus the equation of time. The coordinates of the station are read from a topographical map, accurate to about 0.01, and the Greenwich-time can be determined to about 0.1 sec. The accuracy of the astronomical north depends in the first place on the precision of the time-readings. It is further determined by the quality of the observation, that means, by the accuracy of the levelling during the observation. The quality of the determination will appear from the scattering of the values obtained by the series of determinations A and B , and from the difference $A - B$. The magnitude $(A - B) \cos h$ must have the same value at different stations, as long as the instrument remains unchanged. Finally a good measurement is accurate to 0.1.

If the sky was cloudy, the direction of the astronomical north was determined by the help of triangulated points. The Government's Triangulation (Rijksdriehoeksmeting) and the other survey services have determined the coordinates of a great number of church-steeple, triangulation stones and cadastral points, with respect to a system of coordinates with origin nearly coinciding with O. L. Vrouwe-tower at Amersfoort¹⁾. If the theodolite is placed above a triangulated point with the coordinates x_1, y_1 and is aimed at a spire with the coordinates x_2, y_2 , then the telescope will make an angle arc $\operatorname{tg} (x_1 - x_2) / (y_1 - y_2)$ with the direction of the ordinate through the place of observation. At the origin of the system of coordinates the direction of the y -axis coincides with the direction of the true north; at an arbitrary point with the coordinates xy an angle p exists between the two directions found by the formula²⁾

$$p = 0,04155 x + 1,1 \cdot 10^{-8} xy + \dots$$

x and y must be expressed in meters, and p is found in seconds of arc.

The direction of the ordinate must be corrected for the angle p in order to find the astronomical north.

In cases where the triangulated point cannot be used directly, owing to the surroundings being magnetically disturbed, the theodolite can be placed on the line between this point and a tower.

¹⁾ Rechthoekige coördinaten, Rijksdriehoeksmeting 1855—1928, Delft 1929. Militair register getrianguleerde punten, 1933.

²⁾ Hk. J. HEUVELINK, De stereografische kaartprojectie in hare toepassing, Delft, 1910.

It often happens, that no triangulated stone is present on the spot where the observation must be done. Then it is possible to determine the coordinates of the theodolite by aiming at three triangulated steeples. For this purpose we used a calculating-scheme, which gives results considerably quicker than the usual methods and is more comprehensible.

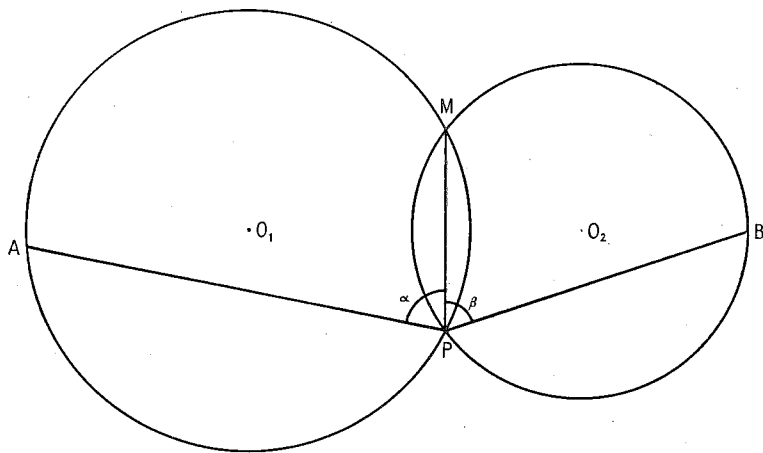


Fig. 2.

Let the triangulated points be A, M and B with coordinates $A(x_1, y_1)$, $M(x_3, y_3)$ and $B(x_2, y_2)$. The angles α and β are measured and it is required to find the coordinates x_p and y_p of the point P. P must be found as the point of intersection of the circles through A, M, P and B, M, P with the angles α and β respectively.

The origin of the system of coordinates is translated to M. The coordinates of a point (xy) are transformed into (s, t) according to:

$$\begin{aligned} s &= x - x_3 \\ t &= y - y_3 \end{aligned}$$

so that the new coordinates of A, M, B and P will be: $A(s_1, t_1)$, $B(s_2, t_2)$, $M(0, 0)$ and $P(s_p, t_p)$.

The coordinates of the centres o_1 and o_2 are:

$$\begin{aligned} s(o_1) &= \frac{1}{2}s_1 + \frac{1}{2}t_1 \cotg \alpha & s(o_2) &= \frac{1}{2}s_2 + \frac{1}{2}t_2 \cotg \beta \\ t(o_1) &= \frac{1}{2}t_1 - \frac{1}{2}s_1 \cotg \alpha & t(o_2) &= \frac{1}{2}t_2 - \frac{1}{2}s_2 \cotg \beta \end{aligned}$$

If A, M and B have arbitrary positions with respect to the point P the signs of α and β are determined by the arcs MA and MB, which must be taken positive if their sense is clockwise with respect to P.

The equation of MP is:

$$t = -s \frac{s(o_2) - s(o_1)}{t(o_2) - t(o_1)}$$

The point of intersection P of the circles o_1 and o_2 with the line PM is given by:

$$\begin{aligned} \frac{1}{2}s_p &= \frac{\{s(o_1)t(o_2) - s(o_2)t(o_1)\} \{t(o_2) - t(o_1)\}}{\{s(o_2) - s(o_1)\}^2 + \{t(o_2) - t(o_1)\}^2} \\ \frac{1}{2}t_p &= \frac{\{s(o_2)t(o_1) - s(o_1)t(o_2)\} \{s(o_2) - s(o_1)\}}{\{s(o_2) - s(o_1)\}^2 + \{t(o_2) - t(o_1)\}^2} \end{aligned}$$

whereas the coordinates of P finally are derived from:

$$\begin{aligned} x_p &= s_p + x_3 \\ y_p &= t_p + y_3 \end{aligned}$$

After determining the astronomical north, either with the help of the sun or by means of a triangulation, the direction of the magnetical north must be found. For this purpose a hollow magnet is suspended in the axis of the theodolite, one end provided with a little lens and the other with a divided scale, which is fitted at the focus of the lens. A silk thread for suspending the magnet is stretched by a weight equal to that of the magnet in order to remove the torsion. This being completed the magnet is attached to the thread and the theodolite is adjusted so that the centre of the scale division coincides with the vertical wire of the telescope.

About five readings are taken, each time just at the second at which a point record is made in the magnetic observatory at Witteveen. Then the magnet is removed from the thread, the tube with the thread is rotated through 180° ; the magnet is turned upside down and attached again. Five other readings are taken, which may all differ some minutes of arc from the first series. If both series are averaged, the eccentricity of the suspending-point as well as the difference between the optical axis and the magnetical one of the telescope will be eliminated. Then the magnet is replaced by a torsion-weight and the angle of torsion of the thread if present can be determined. As the torsion-coefficient of the thread is known, a correction for the value of the astronomical north can be calculated from this angle.

Examples of declination-measurements by observation of the sun as well as by means of triangulation which are worked out completely, can be seen on pages 12 and 13.

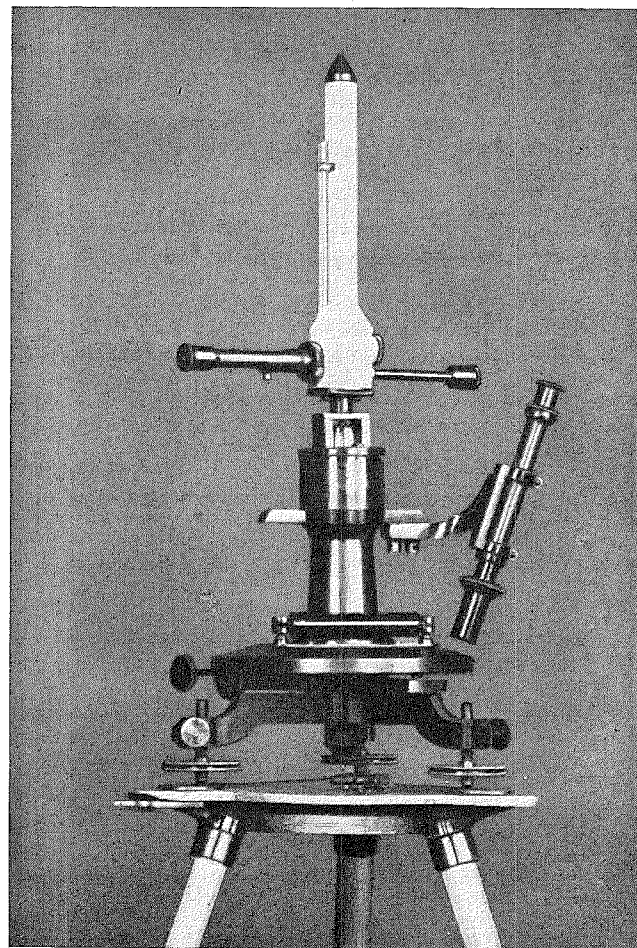


Fig. 3. QHM (quartz horizontal-intensity magnetometer).

Measuring the horizontal component. This component was determined by a QHM (quartz-horizontal-magnetometer) after the design of LA COUR (fig. 3). This instrument is fitted on a magnetic BAMBERG theodolite, left by VAN RIJCKEVORSEL. The QHM contains a little magnet 1.5 cm in length, which is suspended by a quartz-wire. A little mirror is attached to the magnet. The position of the magnet is read by means of a telescope, provided with a GAUSS ocular. When the magnet hangs with its axis in the direction of the axis of the telescope the image of the division scale coincides with the object in the ocular. This coincidence can be accurately adjusted whereupon the corresponding position of the theodolite is read.

Let the quartz-wire, with a torsion-coefficient T , be rotated through an angle β , and let this cause the magnet to make an angle δ with the magnetic meridian, then

$$MH \sin \delta = T\beta$$

In this equation M is the magnetic moment of the magnet, and H is the horizontal component of the earth's magnetic field.

Then the theodolite together with the telescope and the suspending-point of the quartz-wire is rotated through an angle $2\pi + \alpha_1$, till the reflection of the scale division coincides once more. In this case the little magnet, apart from a possible change of H , must have rotated over an angle of exactly 2π . The equation of equilibrium reads:

$$MH \sin (\delta + \alpha_1) = T(\beta + 2\pi)$$

The same process can be carried out by rotating the theodolite in the opposite sense. If this time a rotation of $2\pi + \alpha_2$ appears to be necessary for the coincidence of the reflected image and the scale division itself, it follows that:

$$MH \sin (\delta - \alpha_2) = T(\beta - 2\pi)$$

From this we get:

$$MH \cos \frac{2\delta + \alpha_1 - \alpha_2}{2} \sin \frac{\alpha_1 + \alpha_2}{2} = 2\pi T$$

The instrument is constructed in such a way that in the initial position the angle δ is very small. This angle can be derived from the first three equations.

One finds:

$$\operatorname{tg} \delta = \frac{\sin \alpha_1 - \sin \alpha_2}{2 - (\cos \alpha_1 + \cos \alpha_2)}$$

If δ is indeed small, the angles α_1 and α_2 are almost equal, and one can write:

$$H = \frac{2\pi T}{M \sin \varphi}, \text{ where } \varphi = \frac{\alpha_1 + \alpha_2}{2}$$

φ is the mean deviation of the magnet, caused by a torsion over an angle 2π .

The magnetic moment M and the torsion-coefficient T depend on the temperature t . Besides, the value of M depends on the induction of the magnet caused by the geomagnetic field. Consequently M and T must be written as:

$$T = T_0(1 - c_1 t)$$

$$M = M_0(1 - c_2 t)(1 + \mu H \cos \varphi)$$

where T_0 = the torsion coefficient of the quartz-wire at a temperature of 0°C , M_0 = the moment of the magnet at 0°C in the absence of an external field, c_1 = the tem-

perature coefficient of the quartz-wire, c_2 = the temperature coefficient of the magnet, and μ = the induction coefficient of the magnet.

c_1 , c_2 and μ are small constants. Therefore one can write approximately:

$$\log H = \log \frac{2\pi T_0}{M_0} - \log \sin \varphi + (c_2 - c_1) t \log e - \frac{\mu H \cos \varphi}{\log e}$$

The last term is practically constant over the whole region of the Netherlands. The equation can be simplified to:

$$\log H = c + \alpha t - \log \sin \varphi$$

The constant c depends almost entirely on the torsion coefficient of the quartz-wire and on the magnetic moment of the magnet. This constant, together with the temperature coefficient α , is determined in the Danish Meteorological Institute at Copenhagen, which furnished the instruments QHM 13 and QHM 14. During the measurements the constant c was regularly checked in the magnetic observatory at Witteveen; α was supposed to remain constant, this was not checked by us.

An example of a complete determination is shown on page 14. The reduction of the measurements and the outcome of the calibrations will be described in chapter II.

Measuring the vertical component.

This was done with a BMZ (magnetometric zero balance), after the design of LA COUR (fig. 4). In principle the BMZ is a magnetic balance, constructed in such a way, that the centre of gravity is exactly under the centre of rotation when the magnetic axis of the balance-magnet is

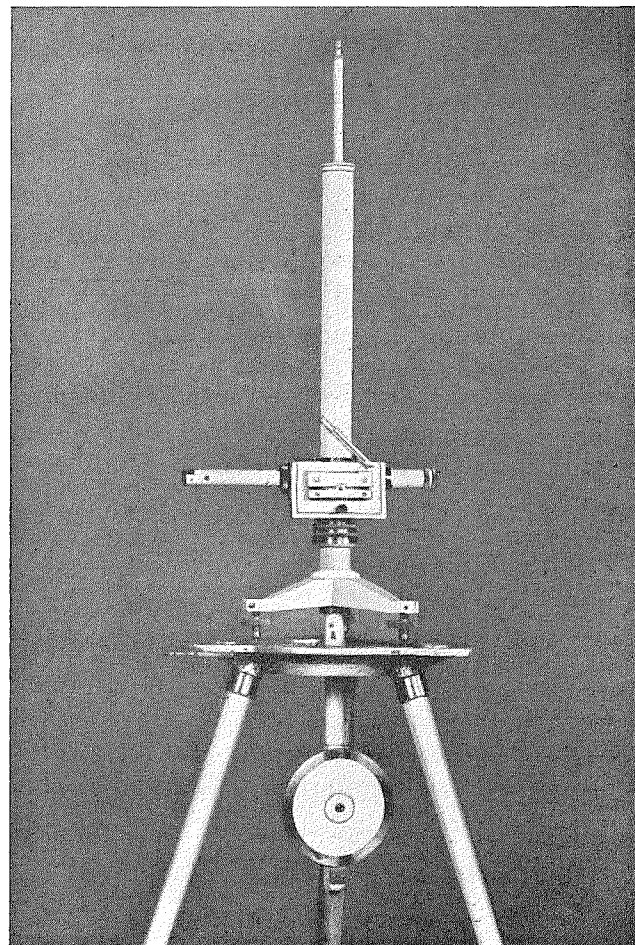


Fig. 4. BMZ (magnetometric zero balance).

horizontal. Over the balance a magnet is fitted, with a magnetic moment neutralizing the vertical component of the geomagnetic field at the place of the balance-magnet. At a distance of 30 cm below the balance a revolving auxiliary magnet is fixed for compensating completely the earth's magnetism. This once being the case, the balance-magnet will have a horizontal position, and this position will be independent of the azimuth of the balance. Whereas the position of the balance generally differs in various azimuths, the neutral position of the balance-magnet will be the same in any direction if the compensation of the vertical component is complete. This neutral position must be determined first of all.

Let the mass of the balance-magnet be P , the magnetic moment M , the distance between the centre of gravity and the rotation-axis a , the azimuth of the magnetic axis δ , the angle between the magnetic axis and the horizontal plane η , and the acceleration of gravity g , then the equation of equilibrium of the balance-magnet will be

$$ZM = (Pga + HM \cos \delta) \operatorname{tg} \eta.$$

The sensitivity of the balance measured in oersted per radian is

$$\frac{dZ}{d\eta} = \frac{Pga + HM \cos \delta}{M \cos^2 \eta}$$

Now P is about 2.5 gram, a is about 10^{-2} cm, M is about 100 c.g.s. units and in this country H is about 0.18 oersted.

Near the neutral position ($\eta = 0$) the sensitivity of the balance in the northern direction ($\delta = 0$) is

$$\frac{dZ}{d\eta} = 0.43 \text{ or } 12\gamma \text{ per minute}$$

and the sensitivity in the southern direction ($\delta = 180^\circ$) is

$$\frac{dZ}{d\eta} = 0.07 \text{ or } 2\gamma \text{ per minute}$$

In the south-position the sensitivity is a maximum, and in this country it is 6 times greater than in the north-position. If the reading of the balance in the south-position is x_2 , and in the north-position x_3 , the unknown neutral position being x_1 , it follows that

$$x_2 - x_1 = 6(x_3 - x_1)$$

from which

$$x_1 = x_3 + \frac{1}{5}(x_2 - x_3)$$

The neutral position is thus found by taking a fifth part of the difference between the readings in the north- and in the south-position, and by adding this to the reading in the north-position, in a direction opposite to the south-reading.

The neutral position being determined in this way, the vertical component Z can be derived directly from the field strengths Z_1 and Z_2 due to the main magnet and the auxiliary magnet.

$$Z = Z_1 + Z_2$$

Let the magnetic moment of the main magnet be M' at a temperature of 0°C , its temperature coefficient α_1 , the distance between the centres of the main magnet and the balance magnet a , and the coefficient of expansion of the BMZ α_2 , then at a temperature of $t^\circ\text{C}$

$$Z_1 = \frac{2M'(1 - \alpha_1 t)}{a^3(1 + \alpha_2 t)^3}$$

Approximately Z_1 can be written as:

$$Z_1 = \frac{2M'}{a^3} \{1 - (\alpha_1 + 3\alpha_2)t\} = Z(1 - \beta t)$$

The field strength Z_2 depends on the angle between the magnetic axis of the auxiliary magnet and the horizontal plane in the first place, and moreover on the temperature. As the two magnets are made of the same magnetic steel

$$Z_2 = Z_p(1 - \beta t)$$

where Z_p can be derived from the dimensions of the auxiliary magnet, the magnetic moment of this magnet at a temperature of 0°C , the distance to the balance magnet, and the angle between the auxiliary magnet and the horizontal plane, after the formulas of AD. SCHMIDT. We obtain then:

$$Z = Z_c + Z_p - (Z_c + Z_p)\beta t$$

which can be simplified to

$$Z = Z_c + Z_p - \alpha t$$

where α is a temperature coefficient, proportional to $Z_c + Z_p$. In this country the variations of the vertical component are small, and therefore α has practically the same value everywhere.

Finally the difference between the temperature of the thermometer and the main magnet must be taken into account, in case the temperature is rising or falling. Suppose that the temperature of the thermometer in the main magnet varies by an amount Δt per minute, a stream of heat will flow from the air via the main magnet to the thermometer or in the opposite sense, so that the temperature of the main magnet will be $t + \varepsilon \Delta t$, ε is a constant depending on the heat insulation of the main magnet. The vertical component is ultimately determined from

$$Z = Z_c + Z_p - \alpha t - \varepsilon \alpha \Delta t$$

We used the BMZ 26. The values of Z_c and a table of Z_p at different angles were determined in the Danish Meteorological Institute. The coefficients α and ε have also been determined there. At fixed times the BMZ was compared with the values of the vertical component in the magnetic observatory, which were derived from the absolute measurements of the horizontal component and the inclination. The coefficients α and ε were supposed to be invariable, and they were not checked. It will appear from the calibrations — see chapter II — that no large variations of Z_c or Z_p were observed.

In the northernmost part of the country, on the Wadden islands, the moment of the revolving auxiliary magnet was too small to compensate the field completely. In this region a fixed auxiliary magnet was used, fitted below the rotating-disc. The field strength was determined by means of the rotating-disc — see next chapter.

CHAPTER II

THE REDUCTION OF THE MEASUREMENTS

The reduction of a declination measurement depends entirely on whether the direction of the astronomical north was determined by observation of the sun or by means of triangulated points. An example of a sun's observation is printed on page 12. At the top to the right the astro-

nomical longitude and latitude of the station are filled in, read from an accurate map. For this purpose we used the topographical map of the Netherlands, scale 1 : 50 000, issued by the Topographical Service. Moreover, the equation of time, valid during the sun's observation, is indicated and the longitude converted into time. The correction for the chronometer is given as well.

The first column of the astronomical part of the measurement contains 20 readings of the moment the sun touches the reticle-wire, with the right limb and with the left one alternately. After every four readings either the mirror or the whole theodolite has been turned. The read times and angles are checked graphically for possible errors. The average of each group of observation times is converted into mean local time (column 2), taking into account the correction to the chronometer and the geographical longitude of the station. The average of the circle readings is corrected in column 4 for errors in the division of the circle. In column 5 the true hour-angle of the sun is computed, belonging to the average of each group of observations. The right-hand part of the example contains the computation of the sun's azimuth α . Subtracting this from the circle reading when the measurement was carried out after true noon, and adding it to the circle reading for an observation before noon, the direction of the instrumental north is obtained. In this way we obtain six numbers, which are combined two by two to give A_1 , B and A_2 , being the values of the instrumental north with the sun in front of the observer, behind him and in front of him again. If the measurement has been carried out well A_1 and A_2 must be equal. By comparing them it can be seen immediately whether a large error has slipped into the data used for the computation. The average of A_1 and A_2 is now combined with B to obtain the value of the astronomical north, at the foot of the form to the right.

Below, the magnetical part of the measurement is shown. The columns 2 and 3 contain the times of observation with the corresponding circle readings. Column 4 contains the readings of the declination from the magnetogram in the magnetic observatory at Witteveen. It is supposed now, that the geomagnetic variations are synchronous and have the same amplitude at Witteveen as at any other place in the Netherlands. This hypothesis, a priori plausible on account of the small distances in our country compared with the distances at which the currents responsible for these variations, flow in general, has been tested by a comparison between the magnetograms of Witteveen and De Bilt. It appeared that the magnetograms are indeed identical for small variations, and that only in case of strong disturbances small differences occur.

The read positions of the theodolite are now reduced to one definite value of the declination at Witteveen, indicated as base D . The corrections, that means the differences between the declination at the moment of the measurement and this basic value, are indicated in column 5 converted into minutes of arc. Column 6 contains the reduced circle readings. The scattering of the values gives an idea of the accuracy of each individual reading. Abnormal values can then be easily rejected. Column 7 gives the mean of the readings with the magnet upright and

in the opposite position. Column 8 contains the value of the torsion of the thread, read after the observation was carried out. As the torsion amounted to 0° at the beginning of the observation, half of the torsion angle is applied as a correction to the obtained values of the magnetical meridian. This is done in the lower part of column 8. Finally, column 9 contains the computation of the declination, the instrumental correction determined by calibrations in the magnetic observatory, and the definite reduced declination.

In the second place fig. 6 shows an example of a declination measurement, the direction of the astronomical north being fixed with the help of triangulated points. In column 1 the six church-steeple are mentioned, visible at the station Wognum. Column 2 contains the corresponding readings of the theodolite, column 3 the corrected readings, column 4 and 5 the coordinates of the spires. The identification of the church-steeple often gave some trouble and could not always be carried out during the measurement. The coordinates of the station can be read approximately from a topographical map, so that the church-steeple can be identified afterwards by the read direction. From a number of six church-steeple three combinations of three steeples are given, which are favourable for the determination of the coordinates of the station. Applying the formulas derived in the last chapter, three values are obtained for these coordinates, which mostly differ a few dm. The average of them is used for the computation of the azimuth of the spires, seen from the station. By adding the value of the azimuth to the circle reading the direction of the y -axis of the coordinates used is obtained. This direction is shown in the last column. Each steeple used gives a value for it. The values for the direction of the y -axis may not differ too much, provided no errors in the coordinates or in the computations have slipped in. The mean value obtained on giving different weights to the azimuth of the various spires, is used for the computation of the astronomical north. For this purpose the correction-angle p is derived (see chapter I) and is added to the direction of the y -axis, giving the value of the astronomical north. The procedure for the magnetical part of the example is the same as in the case of the measurement by observation of the sun.

The reduction of the H-measurement.

In the example of a H -measurement (fig. 7) the instrument used is mentioned in the first column. The third column contains the readings of the theodolite, the fourth column the times read from the chronometer, the fifth column the times reduced to Greenwich-time. Column 6 contains the temperatures, column 7 and 8 the readings of the declination and the horizontal component from the magnetogram of the magnetic observatory at Witteveen. Heading these columns the basic values are given, to which the individual readings are reduced. Column 9 contains the deviations of H from the basic value, column 10 the correction which must be applied to the read position of the circle in order to reduce H to the basic value. It follows namely from the equation for the QHM

$$\log H = c + \alpha t - \log \sin \varphi$$

D-bepaling.

Datum: Woensdag, 28 Augustus 1946.

Plaats: *Veere* $\lambda = 3^{\circ} 38.37'$
 Chronometer: *Nardin* $\lambda = 14^{\circ} 33'$
 Minuutteken van de registrering op 30 sec. Correctie: 12 u G.M.T. = 11-59-50.8 Tijdsvereffening = 1 m 16 s $\lambda = 14^{\circ} 33'$
 Ware tijd = middelh. tijd - tijdsvereffening. Basis D =

Astronomisch gedeelte.

Zon voor	Tijd	G.M.T.	Cirkel	Berekening ϕ	Berekening τ	δ	A ₁	A ₂	B	A ₃
	15 35	15 36	22 20	353 16' 25"	15 50 53	9° 47.2	9.779365	9.779365	9.779365	9.779365
	35	36	53 50	37	16		9.23669	9.23669	9.23669	9.23669
	36	36	53 38							
	37	37	53 10							
Spiegel omgedraaid										
	38	38	59 40	353 50 50	3 49 37	9° 47.2	9.03034	9.03034	9.03034	9.03034
	38	38	59 10				0.10722	0.10722	0.10722	0.10722
	39	39	59 11				9.89386	9.89386	9.89386	9.89386
	39	39	59 41				9.70056	9.70056	9.70056	9.70056
Zon achter										
	40	40	03 10	174 30 42	3 52 12	9° 47.1	9.61746	9.61746	9.61746	9.61746
	41	41	03 00	60	00 00		0.41444	0.41444	0.41444	0.41444
	41	41	03 40				0.10722	0.10722	0.10722	0.10722
	42	42	03 58				0.30719	0.30719	0.30719	0.30719
Spiegel omgedraaid										
	43	43	33 50	175 01 47	15 58 16	9° 47.1	9.92577	9.92577	9.92577	9.92577
	43	43	33 30	60	00 00		9.93200	9.93200	9.93200	9.93200
	44	44	33 48				9.46404	9.46404	9.46404	9.46404
	44	44	33 29				0.45653	0.45653	0.45653	0.45653
Zon voor										
	45	45	41 40	355 28 05	16 00 47	9° 47.0	0.44125	0.44125	0.44125	0.44125
	46	46	41 30	39	16 00 47		70° 05' 52"	70° 05' 52"	70° 05' 52"	70° 05' 52"
Spiegel omgedraaid										
	47	47	59 00	355 50 50	3 57 32	9° 47.0	353 17 02	353 17 02	353 17 02	353 17 02
	48	48	59 41				283 47 58	283 47 58	283 47 58	283 47 58
	48	48	59 00				283 46 00	283 46 00	283 46 00	283 46 00
	49	49	59 11				0.29163	0.29163	0.29163	0.29163
	49	49	59 14				9.93700	9.93700	9.93700	9.93700
	49	49	59 29				9.45608	9.45608	9.45608	9.45608
	49	49	59 51				0.46996	0.46996	0.46996	0.46996
	49	49	59 58				0.49033	0.49033	0.49033	0.49033
	49	49	59 58				71° 42' 53"	71° 42' 53"	71° 42' 53"	71° 42' 53"
	49	49	59 58				355 28 44	355 28 44	355 28 44	355 28 44
	49	49	59 58				283 45 57	283 45 57	283 45 57	283 45 57
	49	49	59 58				283 45 48	283 45 48	283 45 48	283 45 48
	49	49	59 58				283 45 48	283 45 48	283 45 48	283 45 48

Magnetisch gedeelte.

Stand	Tijd	G.M.T.	Cirkel	D	ΔD	Gemiddeld	Torsie	Astronomisch Noord	Magnetisch Noord	Declinatie	Instrument correctie	Herleide declinatie
	15 05	15 05	275	90.7	-19.6	74° 01' 11"	+5"	283° 45' 0"	275	8° 10.9	-	8 17.0
	07	07	280	90.3	-19.7			275 26.9				
	08	08	284	90.2	-19.7							
	09	09	280	90.8	-19.8							
	10	10	279	90.8	-19.9							
	17	17	28.2	90.0	-19.9							
	18	18	28.5	90.0	-20.0							
	19	19	28.5	89.9	-20.1							
	20	20	28.8	89.8	-20.2							
	21	21	29.1	89.7	-20.2							
	15	15	295	89.7	-20.2							

Fig. 5. Measurement of declination by sun's observation.

D-bepaling.

Plaats: *Wegman*
Chronometer: *Nardin*

Minuutteken van de registrering op 30 sec

Datum: *Donderdag dag, 10 Augustus 1944*

Correctie: 12 u G.M.T. = 12-00-18,0

$\varphi = 52^{\circ} 41,02'$ $\lambda = 5^{\circ} 0,47'$
Tijdsvereffening = $\frac{m}{s}$
Basis D = 63,9 $\frac{m}{s}$ Ware tijd = middelb. tijd — tijdsvereffening.
(= 6° 43,0 *Wittereen*)

Astronomisch gedeelte.

Zon voor	Tijd	G.M.T.	Cirkel	Berekening ϕ	Berekening τ	δ	A ₁	B	A ₂
Spiegel omgedraaid									
1. H.K.	105' 39" 30	105' 30"	105' 40"	-23670.48	+60846.33		1.	105' 40" 40"	60 34' 45"
2. R.K.	54' 37" 10	54' 37"	54' 37"	-26170.00	+63143.86		2.	54' 37" 53"	60 34' 43"
3. R.K.	8' 00" 00	8' 00"	8' 00" 45"	-30052.98	+62124.15		3.	8' 00" 45"	60 34' 42"
4. H.K.	8' 40" 00	8' 40"	40' 45"	-30600.47	+62634.64		4.	8' 40" 45"	60 34' 30"
5. R.K.	359' 23" 10	359' 23"	359' 23" 52"	-28687.09	+60439.25		5.	359' 23" 52"	60 34' 45"
6. H.K.	351' 58" 00	351' 58"	351' 58" 36"	-29990.54	+60479.61		6.	351' 58" 36"	60 34' 58"
Spiegel omgedraaid									
Zon voor									
			Combinatie						
			I (1,3,2)	-25718.12	+58805.80				60 34' 42"
			II (1,4,2)	-25718.60	+58805.77				- 17 36
			III (1,6,2)	-25717.52	+58805.73				60 34' 42"
			Spieiddekl	-25718.08	+58805.77				
Spiegel omgedraaid									

log cos ϕ =	
log tg δ =	
log B =	
log sin ϕ =	
log cos τ =	
log A =	
A =	
B =	
A-B =	
log sin τ =	
log(A-B) =	
log tg α =	
α =	
Cirkel =	
Instr. Nrd =	
Zonhoogte h =	
cos h =	

Magnetisch gedeelte.

Stand	Tijd	G.M.T.	Cirkel	D	ΔD	Gemiddeld	Torsie
0	20' 48"	6	35	51.6	-12.3	233 24.0'	
	21	21	35	51.6	-12.2	233 24.0'	-40"
	22	22	35	51.7	-12.2	233 24.0'	
	23	23	35	51.7	-12.2	233 24.0'	
	24	24	35	51.6	-12.2	233 24.0'	
1	27	27	32	51.8	-12.0	233 24.4'	
	28	28	32	51.9	-12.0	233 24.7	
	29	29	32	51.9	-11.9	233 24.7	
	30	30	33	52.0	-11.9	233 24.7	
	31	31	32	52.0	-11.8	233 24.4'	
	30	6	32	52.0	-11.8	233 24.4'	
						Correctie voor torsie	
						$\frac{-20}{360} \times 9' =$	
						-0.5	

Fig. 6. Measurement of declination by means of triangulated points.

Station: 376
 Plaats: Slenaken
 Chronometer: Alpina

H-bepaling Datum Dinsdag, 7 Sept. 1943

$\varphi = 50^{\circ} 46.96' \lambda = 5^{\circ} 50.62'$

Correctie: 12u G.M.T. = 14 u 00 m 0 s s kloktijd.

Minuutteken van de registrering op 30 sec. Basis H: 80.0' mm (= 85.9' mm μ gr = 17960 Wittenveen)

Toestel	Stand	Cirkel		Tljd		G.M.T.		T _B = 24.0	Temp.	D	H	ΔH_{mm}	$\Delta H'$	$\Delta D'$	$\Delta \rho_T$	Cirkel	Gemidd.	Opm.	
		0	1	h	m	s	h												m
QHM14	hul	222	87	16	03	33	14	03	30	66.0	79.2	-0.3	-0.6	+1.0	-0.7	203	25.3	222 09.7	$\varphi = 61^{\circ} 16.7'$ $\rho = 24.5^{\circ} C$ $H = 79.5$ mm
		283	156	"	07	"	"	07	"	66.0	79.2	-0.3	-0.6	+0.9	-0.9	25.6	25.6		
		283	261	"	08	"	"	08	"	65.9	79.2	-0.3	-0.6	+0.9	-0.9	25.3	25.3		
		283	259	"	09	"	"	09	"	65.9	79.2	-0.2	-0.5	+1.0	-0.9	25.2	25.2		
		283	258	"	10	"	"	10	"	66.0	79.2	+0.1	+0.3	+1.0	-0.2	27.0	27.0		
		283	260	"	11	"	"	11	"	66.0	79.2	+0.4	-0.8	+0.7	0.0	51.7	51.7		
QHM13	hul	352	260	16	36	33	14	36	30	65.0	81.0	-0.5	-1.1	0.0	-0.2	53	26.2	352 26.0	$\varphi = 61^{\circ} 00.4'$ $\rho = 22.7^{\circ} C$ $H = 81.0$ mm
		53	273	"	41	"	"	41	"	65.3	80.4	-0.1	-0.2	+0.4	0.0	27.2	27.2		
		53	270	"	42	"	"	42	"	65.4	80.9	0.0	0.0	+0.5	-0.2	26.9	26.9		
		53	270	"	43	"	"	43	"	65.5	81.0	-0.2	-0.4	+0.5	-0.2	27.2	27.2		
		53	270	"	44	"	"	44	"	65.5	80.8	-0.1	-0.2	+0.6	-0.2	26.9	26.9		
		53	271	"	45	"	"	45	"	65.6	80.9	-0.2	-0.2	+0.2	-0.2	25.2	25.2		
QHM14	rechts	291	247	"	50	"	"	50	"	65.2	80.7	-0.2	+0.5	+0.2	-0.2	25.2	25.2		$\varphi = 61^{\circ} 00.4'$ $\rho = 22.7^{\circ} C$ $H = 81.0$ mm
		291	252	"	51	"	"	51	"	65.2	80.8	-0.2	+0.4	+0.2	-0.2	25.6	25.6		
		291	207	"	52	"	"	52	"	65.2	80.8	-0.2	+0.4	+0.2	-0.2	26.1	26.1		
		291	261	"	53	"	"	53	"	65.1	80.8	-0.2	+0.4	+0.1	-0.2	26.4	26.4		
		291	264	"	54	"	"	54	"	65.1	80.9	-0.1	+0.2	+0.1	-0.2	26.7	26.7		
		291	264	"	54	"	"	54	"	65.1	80.9	-0.1	+0.2	+0.1	-0.2	26.7	26.7		

Q H M 13
 $c = 9.21245 - 10$
 $d = 0.000163$
 Q H M 14
 $c = 9.21276 - 10$
 $d = 0.000159$

Eerste meting
 $c = 9.21276 - 10$
 $d = 0.00381$

Tweede meting
 $c = 9.21245 - 10$
 $d = 0.00370$

Gemiddeld:
 $H = 18793.5$

Eerste meting:
 $H = 18790$
 Basis Corr. = $+3$

$\log \sin \varphi = 9.21615$
 $\log H = 9.94185$
 $\log H = 0.27430 - 1$
 $\text{Instr. Corr.} = 18806$
 $\text{Instr. Corr.} = -7$
 18799

Tweede meting:
 $H = 18793$
 Basis Corr. = -5

Tweede meting:
 $H = 18799$
 Basis Corr. = -5

Fig. 7. Measurement of the horizontal component by means of QHM 13 and 14.

de Bief

Z-meting met B M Z 26.

Station Nr.: 194

Datum: Zou dag, 14 Juli 1946

$\varphi = 52^{\circ} 48' 8'' \lambda = 6^{\circ} 40' 1''$ Opmerkingen: weiland achter de parolpen
 Minuutreeken op 30 sec.
 Basis Z: 29.4 mm = 43854 BMZ
 1 mm Z = 5.40 f

Chronometer: Alpina
 Nulpuntcorrectie p = -0.20
 normale positie = 30.9

Correctie: 12 u G.M.T. = 13-00-00
 Hoofdcostante H = 43401 γ
 Z = 43401 + Zp - at - 1,5 Δ t.

Klok tijd	G.M.T.		Draaischijf	D-p	t	Δt	Z	Z Witveen	ΔZ	Herleide Z	Gemiddeld					
	43401	43401														
5	03	30	4	03	30	66°	95	66°	75	20	98	43536	43853	-113	43423	
"	04	"	"	04	"	66	95	66	95	20	80	35	"	"	22	
"	05	"	"	05	"	67	30	67	10	20	58	36	"	"	23	
"	06	"	"	06	"	67	50	67	30	20	35	36	"	"	23	43424
"	07	"	"	07	"	67	55	67	35	20	15	38	"	"	25	
"	09	"	"	09	"	67	80	67	60	19	70	40	"	"	27	
"	17	"	"	17	"	68	60	68	40	18	48	43543	43853	-113	43430	
"	18	"	"	18	"	68	55	68	35	18	35	45	"	"	32	
"	19	"	"	19	"	68	60	68	40	18	25	45	"	"	32	
"	20	"	"	20	"	68	70	68	50	18	15	45	"	"	32	43431
"	23	"	"	23	"	69	00	69	80	17	90	44	"	"	31	
"	30	"	"	30	"	69	15	69	95	17	60	43543	43853	-113	43430	
"	31	"	"	31	"	69	15	69	95	17	59	43	"	"	30	
"	32	"	"	32	"	69	10	69	90	17	58	44	"	"	31	43430
"	33	"	"	33	"	69	15	69	95	17	52	45	"	"	32	
"	39	"	"	39	"	69	10	69	90	17	45	43547	43853	-113	43433	
"	40	"	"	40	"	69	10	69	90	17	43	47	"	"	33	
"	41	"	"	41	"	69	15	69	95	17	42	46	"	"	32	43433
"	42	"	"	42	"	69	10	69	90	17	41	47	"	"	33	
5	43	30	4	43	30	69	10	69	90	17	42	46	43854	-114	43422	

H = 43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401	43401
Zp = 452+	440+	445+	442+	442+	422+	422+	422+	422+	422+	422+	411+	411+	411+	411+	411+	411+	411+
850	849	846	843	837	825	825	825	825	825	825	812	812	812	812	812	812	812
at = 322-	318-	315-	312-	309-	302-	283-	281-	272-	272-	272-	270	269	269	269	269	269	269
531'	531'	531'	531'	535	535	539	542	542	542	541	543	543	544	544	546	546	547
1,5 Δ t = 4	4	4	5	5	5	3	3	3	3	3	0	0	0	0	0	0	0
Z = 536	535	536	536	538	540	543	545	545	544	544	543	543	544	544	547	546	547

Fig. 8. Measurement of the vertical component by means of BMZ 26.

K 2872

that with a variation of H a variation of φ is related according to

$$dH = -H \cotg \varphi d\varphi$$

Column 11 contains the deviations of D from the basic value. The amounts of these deviations are applied as corrections to the readings, since the magnet, even in its deflected position, will follow the slight rotations of the geomagnetic field during the measurement. Column 12 contains corrections for the reduction of the temperatures to a basic temperature, by means of the relation

$$dH = H\alpha dt$$

Column 13 contains the circle readings reduced to definite values of t , D and H . The mean values of each group of readings are given in column 14. The mean angle of deflection caused by a torsion over an angle 2π is shown in column 15. At the bottom of the example the values of H have been computed. An instrumental correction is added to them, which is determined by calibrating the instrument in the magnetic observatory viz. -7γ for QHM 13 and $+14\gamma$ for QHM 14. Finally a basic correction is given, obtained by reducing a provisional basic correction H in column 8 to the definitive basic value, written at the head of the example. The final results of the two measurements, which were carried out at a distance of 50 m, lead to figures differing only by 1γ .

The reduction of the Z -measurement.

Fig. 8 shows a complete Z -measurement. In the first place a few constants are given viz. the zero point correction of the rotating-disc, determined by calibration in the magnetic observatory, the neutral position determined during the measurement, the temperature coefficient α

given by the Danish Meteorological Institute, and the basic value to which all readings are reduced. Column 1 contains the time of each reading, column 2 the Greenwich time, column 3 the reading of the rotating-disc with the auxiliary magnet, column 4 the corrected reading of this disc, column 5 the temperature of the main magnet, column 6 the temperature gradient per minute, column 7 the value of the vertical component derived from the preceding data, column 8 the value of Z at Witteveen, column 9 the difference between this value and the basic value to which all measurements were reduced, column 10 the reduced Z -values, column 11 the average of the series of Z -values. The BMZ was used at four different places some tens of meters apart, so that four different values of Z are obtained. The last three of them are in good agreement with each other, whereas the first one is different. For that reason only the last three values have been used for the determination of the definitive mean value, the first one being rejected.

The determination of the time.

For reading the time during the astronomical measurements the chronometer Nardin No. 10658 was mostly used. This chronometer was compared with a radio time signal once a day at least. In the years 1942 and 1943 we used the time signals of the Deutsche Seewarte, later on mostly the six pips of the BBC.

Fig. 9 shows the rate of the Nardin No. 10658 during five successive days. It turns out that the rate is constant or nearly constant for a rather long period, so that the time correction between the time signals could be determined by interpolation, with an accuracy of 0.1 sec.

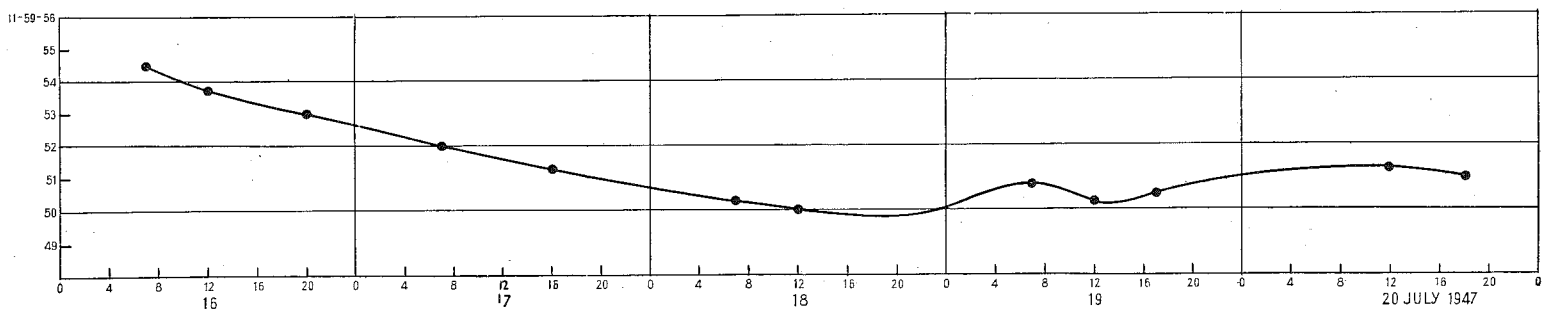


Fig. 9. Rate of the chronometer Nardin 10658 during five successive days.

The calibration of the instruments.

All measurements were reduced to the annual means for the epoch 1945.0, as determined with the absolute instruments in the magnetic observatory at Witteveen. Consequently the instruments for the field work were compared from time to time with these instruments and with the outcome of the absolute measurements performed by the officer-in-charge, who is responsible for the magnetic observatory.

In the years before 1945 the absolute measurements have been performed regularly, and therefore the calibration of the instruments, used for the magnetic survey, is known quite accurately.

The same does not hold however for the years 1945—1948. After the completion of the present publication it has become evident that the absolute measurements made by

the former officer-in-charge of the Witteveen magnetic observatory cannot be considered as trustworthy during these years ¹⁾. Without going into details of this regrettable matter it can be said that the reduction of the measurements was based upon a more or less fictitious value, which was chosen in such a manner that the outcome of the measurements fitted the results of the years 1942—1945.

Whereas it is a matter for regret that accurate calibrations are lacking for the years 1945—1948, the method followed by the former officer-in-charge implies that the field measurements of these years are reasonably trustworthy, as part of them overlapped the measurements of former years. Moreover the greater part of the country was surveyed during the first three years, so that only the measure-

¹⁾ See Preface of Yearbook B, Geomagnetism, 1947, Royal Netherlands Meteorological Institute, De Bilt.

ments of Zeeland, Zeeuws-Vlaanderen, the Zuid-Hollandse islands and the Wadden Sea, fall into the less accurate period.

In order to clear up the uncertainty a couple of measurements was repeated in the above mentioned regions during the fall of 1949 and the spring of 1950, with the result that in the writer's opinion all values in these regions have an accuracy better than 1' and 10 γ for D , H and Z resp.

The D -instrument was compared with the DOVER instrument of the Witteveen observatory. The difference between both instruments was fairly constant. Corrections ranging from $-1'.3$ to $-1'.7$ were applied to all declination values found by means of the field instruments.

For the instruments QHM 13 and QHM 14 the following formulas were used:

$$\text{QHM 13 : } \log H = 9.21245 - 10 + 0.000163 t - \log \sin \varphi$$

$$\text{QHM 14 : } \log H = 9.21276 - 10 + 0.000159 t - \log \sin \varphi$$

By comparing the QHM's with the H -instrument DOVER in the magnetic observatory Witteveen corrections of $+3 \gamma$ and $+2 \gamma$ were found for QHM 13 and QHM 14 respectively in the year 1942. In the beginning of 1943 the quartz-wire of QHM 14 broke during transport. After that the correction was $+14 \gamma$, whereas QHM 13 remained practically constant. In the year 1944 the quartz-wire of QHM 13 broke, and after the repair the correction was -1γ . After the end of 1944 QHM 13 was no longer used for field work but was kept as a reserve instrument in the magnetic observatory at Witteveen. Calibrations during the summer of 1949 led to a correction of -12γ for QHM 14.

The formula used for BMZ 26 was

$$Z = 43401 + Z_p - \alpha t - 1.5 \alpha \Delta t$$

Z_p is the field-strength of the turning magnet, varying between 0 and 1160 γ .

The values of Z_p depend on the position of the turning magnet, and had been calibrated at the Meteorological Institute, Charlottenlund, Denmark. The calibration of BMZ 26 consisted of comparing the instrument with the basic values, which were derived for the vertical component of the geomagnetic field at the Witteveen observatory calcu-

lated from the H -measurement with the DOVER-instrument and from the I -measurements with the earth-inductor TOEPFER. The correction in the year 1942 and in the beginning of 1943 was -9γ . After this date the behaviour of BMZ 26 was very irregular. The scattering between the individual readings was very large, the corrections amounted from -20γ to -60γ . The error appeared to be caused by the balance magnet touching the damping. The magnet moves between copper plates whose purpose is to damp the motion considerably. It turned out that slight displacements of the balance were possible when clamped, from which this touching now and then resulted.

In 1943 the clamping of the balance was corrected, so that it rests now in four forks. When the balance is lowered, it comes in the right position with respect to the damping. After these repairs the behaviour of the BMZ was completely regular and the corrections had once more the former value of -9γ .

New calibrations were made in the year 1949. The whole instrument was placed in a Helmholtz field, generated by three coils, diameter 100 cm, with distances of 40 cm. The outer coils had 50 windings, the middle coil 31, so that the field had a constant value of 90.56 γ per mA along the major part of the axis of the instrument.

It turned out that the magnetic moment of the rotating magnet had increased since 1942 so that all Z_p values had to be corrected by an increase of 1.5 %. The correction for the whole instrument appeared to be -25γ , which can be ascribed to a small change of the magnetic moment of the upper magnet.

The zero point correction of the rotating-disc amounted constantly to $+0^\circ.20$ during these years. It was determined from a comparison of the readings of the rotating-disc in the two positions of the rotating magnet in which the same angle is made with the vertical.

During the measurement in the northernmost part of the country it was necessary to use supplementary magnets for augmenting the compensating field strength. They can be screwed on below the rotating magnet. Their influence upon the balance was determined in 1947 and amounted

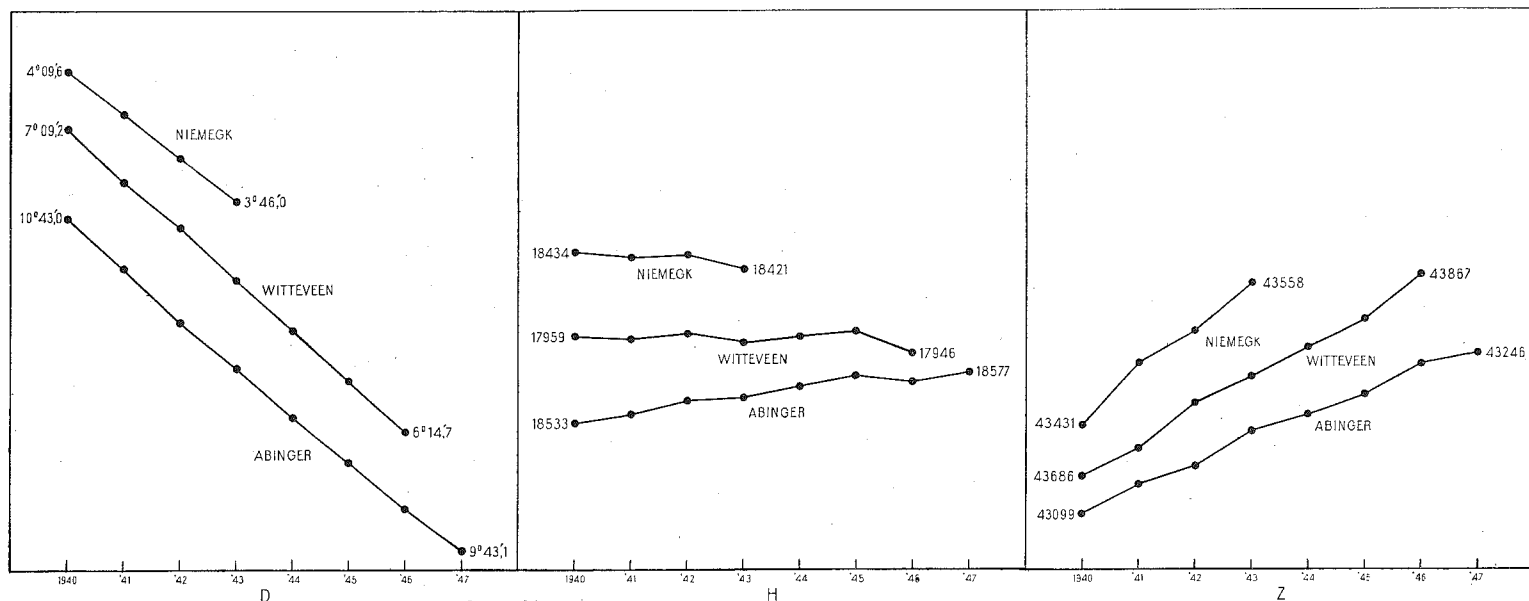


Fig. 10. Secular changes at Witteveen, Abinger and Niemegek.

to 1983 γ for the auxiliary magnet 26-1, to 1988 γ for 26-2, and to 1987 γ for 26-3.

New determinations in 1949 and 1950 gave the values 2005 γ for the auxiliary magnet 26-1 and 2020 γ for 26-3. The magnet 26-2 appeared to be loose in the casing, so that no accurate values could be obtained.

The secular change.

All determinations were reduced to the epoch 1945.0 by applying the secular changes at Witteveen. It was supposed that the secular change has the same value at Witteveen as at any place in the Netherlands. This is certainly allowed as regards the declination. The secular change of D is practically the same everywhere in NW-Europe. However the secular changes of H and Z at De Bilt—Witteveen, Greenwich—Abinger and Potsdam—Niemegk are different.

Below, the annual means of the magnetic elements at the observatories Witteveen, Abinger and Niemegk are given for the recent years. The secular changes are represented graphically in fig. 10.

Annual means of the magnetic elements at Witteveen

$$(\varphi = +52^{\circ} 48'.8; \lambda = +6^{\circ} 40'.1)$$

Year	D	H	Z
1940	7° 09.2	17959	43686
1941	6° 59.8	956	711
1942	51.2	962	750
1943	42.2	954	776
1944	33.0	961	800
1945	24.3	966	827
1946	14.7	946	867

Annual means of the magnetic elements at Abinger

$$(\varphi = +51^{\circ} 11'.1; \lambda = +0^{\circ} 23'.2)$$

Year	D	H	Z
1940	10° 43.0	18533	43099
1941	33.8	539	128
1942	24.8	554	146
1943	16.2	556	172
1944	07.8	566	189
1945	9° 59.5	573	207
1946	51.1	569	235
1947	43.1	577	246

Annual means of the magnetic elements at Niemegk

$$(\varphi = +52^{\circ} 04'.3; \lambda = +12^{\circ} 40'.5)$$

Year	D	H	Z
1940	4° 09.6	18434	43431
1941	01.4	430	484
1942	3° 53.8	433	516
1943	46.0	421	558

To reduce the measurements to 1945.0 the correction for the secular change was applied over a maximum period of 2 years. During this time the difference in the secular change between Potsdam and Witteveen and between Witteveen and Abinger amounted to about 15 γ for H and to about 20 γ for Z . Supposing a linear change of the secular variation with the distance to Witteveen, it appears that in the extreme regions of this country the elements H

and Z showed a variation during the mentioned period, which differs at most 6 γ and 8 γ resp. from the variation at Witteveen.

It is not possible, however, to determine this correction exactly because the annual means at Potsdam are missing after 1943. Besides, for the greater part of the stations this correction is smaller than a few γ , so that it is not taken into account for the present.

The reduced results.

The results of the measurements at 392 stations, reduced to the epoch 1945.0, are tabulated on page 26 till page 30. They contain number, name and coordinates of the stations, together with the values of D, H and Z . In addition, the components in northward direction X and in eastward direction Y are given. These were computed from the formulas

$$X = H \cos D \text{ and } Y = H \sin D$$

All Y -values are negative owing to the fact that in the whole country the declination is westerly.

The last columns contain the values $\Delta X, \Delta Y$ and ΔZ , the deviations of X, Y and Z from a normal field, which is derived in the next chapter. The formulas are:

$$X_n = 18120 - 470\Delta\varphi + 45\Delta\lambda$$

$$Y_n = -2315 + 30\Delta\varphi + 150\Delta\lambda$$

$$Z_n = 43448 + 500\Delta\varphi$$

$$\Delta X = X - X_n; \Delta Y = Y - Y_n; \Delta Z = Z - Z_n.$$

The last 14 measurements (A—N) were carried out after a great part of this publication had already been completed. It was not possible to insert them in the tables; therefore they are printed at the end of the list. At these stations only H and Z have been determined. The values of D , necessary for the calculation of X and Y , were read from the D -chart. These values are therefore given in parentheses.

CHAPTER III

THE CONSTRUCTION OF THE CHARTS

Charts have been constructed of the magnetic elements D, H, Z, X, Y , and of the anomalies $\Delta X, \Delta Y, \Delta Z$. Each chart has been printed in three colours: the numbers of the stations are grey, the values of the magnetic elements are black, and the isolines drawn through them are red.

Although some 240 years have elapsed since HALLEY published his famous isogonic chart, no clear principles for the construction of magnetic charts have been laid down. The tracing of contours is still a matter of individual judgment. A survey of methods of constructing magnetic charts has been given by BERNSTEIN¹⁾.

Two general procedures have been employed in the tracing of contours. They may be termed the analytic method, by which the contours are approximated by a mathematical expression, and the graphical method, consisting of free-hand smoothing. As the network of our stations is dense, and the quality of the measurements is good in general, it was possible to use the graphical method, interpolating the isomagnetic lines between the measured values.

¹⁾ A. BERNSTEIN, A survey of methods of constructing magnetic charts, Terr. Mag. 49, 169, 1944.

Fitting the isolines to the Belgian and German charts.

We have tried to continue the isolines over Belgium and Germany. For Belgium the maps of DEHALU and MARIE MERKEN and HOGE¹⁾ were used, which are reduced to 1913.0. In applying the magnetic difference²⁾ between De Bilt and Witteveen to the mean values of the magnetic elements in the years 1912 and 1913 at De Bilt³⁾, one finds for the magnetic elements at Witteveen for the epoch 1913.0: $D = 12^{\circ} 01'.9$, $H = 18266\gamma$, $Z = 43485\gamma$. The secular changes from 1913.0 to 1945.0 are: $D = -5^{\circ} 33'.2$, $H = -303\gamma$, $Z = +328\gamma$. After applying these values to the Belgian maps the connection between the isogonics is good. Around South-Limburg, however, the network of the Belgian measurements is not dense enough, so that the isogonics could not be continued with sufficient certainty. The connection of the H -isodynamics is good, provided a reduction of -250γ in stead of -303γ is applied. Considering that the Belgian H -measurements as well as the Z -measurements are apparently only accurate to about 10γ , the agreement can not be called bad. The Z -isodynamics too can be connected properly, taking into account the smaller accuracy of the Belgian values.

For the fitting in of the German declination-measurements BURMEISTER's chart⁴⁾ for the epoch 1940.0 was used. According to the annual means at Witteveen the correction for the secular change of the declination from 1940.0 to 1945.0 amounts to $-45'$. After application of this correction the agreement between BURMEISTER's isogonics and ours is rather bad, however. Apparently the borderland of BURMEISTER's chart is based on VAN RIJCKEVORSEL's measurements, which in this region are $5'$ or $10'$ too small. The isolines of our D -chart are therefore traced as well as possible according to the general shape of the German isogonics. Extrapolating the H - and Z -isodynamics was not possible, because the results of the German survey of the second order⁵⁾, which was begun in 1937, were not available.

The construction of the isanomalic chart.

For drawing the chart of anomalies of the magnetic field the knowledge of the normal field is necessary. This normal field is in a certain sense an arbitrary conception, the definition depends entirely on what one wishes to see. If one wants to investigate the continental anomalies over the whole world, one has to assume for a normal field, the field of a magnetic dipole situated in or about the centre of the earth. If, however, one intends to study the regional anomalies in a limited region, the normal field must be chosen in such a way, that it corresponds to the field in the undisturbed parts of this region. One attempts then to write down the components of this normal field as functions of latitude and longitude in an expression of the form

¹⁾ M. DEHALU et MARIE MERKEN, Nouvelle carte magnétique de la Belgique, 1913; E. HOGE, Nouvelle contribution à la carte magnétique de la Belgique, 1934.

²⁾ Yearbook B, Geomagnetism 1945, De Bilt.

³⁾ Annuaire B, Magnétisme Terrestre 1938, De Bilt.

⁴⁾ F. BURMEISTER, Karte der Erdmagnetischen Missweisung für die Epoche 1940.0.

⁵⁾ R. BOCK, Über die magnetische Reichsvermessung II. Ordnung und ihre ersten vorläufigen Ergebnisse, Z.f. Geophysik, 15, 66, 1939.

$$F = F_0 + a\varphi + b\lambda + c\varphi^2 + d\lambda^2 + e\varphi\lambda + \dots$$

F_0 and the coefficients a, b, c, d, e, \dots are to be determined so that the undisturbed field is represented by this expression. It would be possible to calculate these coefficients from the values of the true magnetic field by the method of least squares, if the deviations of the true field from the normal field were distributed according to the laws of chance. This method has indeed been used by SCHOTT⁶⁾ and HAZARD⁷⁾ in constructing the magnetic maps of the United States.

As however was remarked by RÜCKER en THORPE⁸⁾, the anomalies are usually not distributed according to chance. The latter authors followed therefore a much simpler procedure. In the first place they chose for the normal field the expression

$$F = F_0 + a\varphi + b\lambda + e\varphi\lambda + f \cos d(\varphi - g)$$

with a, b, d, e, f and g as unknown factors, as this function was nearly in harmony with the mean course of the isolines. The 205 stations were grouped into a number of districts, and for every district the arithmetical average of the values of latitude, longitude and magnetic element was determined. The unknown coefficients were finally determined from a number of equations, equal to the number of unknowns. This method was also used by VAN RIJCKEVORSEL⁹⁾. He was rightly of opinion, that for a small country as the Netherlands a linear approximation of the normal field must be sufficient. A magnetic centre was determined by taking for all the stations the arithmetical means of the longitude, the latitude and the magnetic components. In each place the deviation of the normal field from that magnetic centre is given by

$$\Delta F = a\Delta\varphi + b\Delta\lambda$$

where for ΔF is successively taken ΔX , ΔY and ΔZ ; $\Delta\varphi$ and $\Delta\lambda$ are the differences between the latitude and longitude of the place in question and the coordinates of the magnetic centre. For each component VAN RIJCKEVORSEL obtained two equations by summing up the equations for all the stations north of the magnetic centre and by doing the same for all the stations east of this centre. The stations in the southwestern quadrant were thus left out of consideration, because the disturbances here were greater than in the rest of the country. The ultimate equations were:

$$\begin{aligned} \Delta X &= -354.1\Delta\varphi + 116.0\Delta\lambda \\ \Delta Y &= -73.0\Delta\varphi - 143.4\Delta\lambda \\ \Delta Z &= +429.8\Delta\varphi - 73.5\Delta\lambda \end{aligned}$$

with the magnetic centre: $\varphi = 52^{\circ} 16' 41''$ N.L., $\lambda = 5^{\circ} 30' 43''$ E.L., and $X = 17706\gamma$, $Y = -4611\gamma$, $Z = 43994\gamma$.

A still simpler method for the construction of the normal field consists of drawing smooth isolines in one way or another, for example, by drawing by free hand or by

⁶⁾ C. A. SCHOTT, Distribution of the magnetic declination in Alaska and adjacent waters for the year 1895, U. S. Coast and Geodetic Survey Bull. 34, 129, 1895.

⁷⁾ D. L. HAZARD, United States magnetic tables and charts for 1925, U. S. Coast and Geodetic Survey, Serial 453, 1929.

⁸⁾ A. W. RÜCKER and T. E. THORPE, Magnetic survey of Great Britain and Ireland, Phil. Trans. 190 A, 53, 1890.

⁹⁾ E. VAN RIJCKEVORSEL l.c.



averaging groups of observations, the isolines being approximated by a linear or quadratic equation. This method is followed by LEWIS¹⁾ in constructing an isogonic map of South-Africa. A combination of the above described methods was used by LJUNGDAHL²⁾. From the measurements of a number as relatively undisturbed selected stations he derived by the method of least squares a set of linear equations for all the magnetic elements.

The formulas for the normal field in the Netherlands, used by HARTMANN³⁾, are based on the geomagnetic field at the observatories at Abinger, Eskdalemuir, Witteveen, Chambon-la-Forêt, Rude Skov and Niemegek. It is supposed that the field at these observatories is determined by the dipole-field plus the continental European anomaly and that, therefore, the regional disturbances are small. This is an arbitrary supposition, though it may be expected that a magnetic observatory is generally built in a magnetically undisturbed region. HARTMANN finds:

$$\begin{aligned} X_n &= 18105 - 470\Delta\varphi + 45\Delta\lambda \\ Y_n &= -2430 + 30\Delta\varphi + 150\Delta\lambda \\ Z_n &= 43375 + 500\Delta\varphi \end{aligned}$$

where $\Delta\varphi = \varphi - 52^\circ$ and $\Delta\lambda = \lambda - 5^\circ$. It turned out, that these formulas gave a reasonable distribution of positive and negative regions for ΔZ , but that the horizontal disturbance-components deviated systematically from the direction, which could result from the vertical disturbance-field. A much better result was obtained by slightly altering the constant term in the formulas for X_n and Y_n . Finally HARTMANN used the following formulas:

$$\begin{aligned} X_n &= 18115 - 470\Delta\varphi + 45\Delta\lambda \\ Y_n &= -2380 + 30\Delta\varphi + 150\Delta\lambda \\ Z_n &= 43775 + 500\Delta\varphi \end{aligned}$$

In this report the anomalies, which are reproduced in the Δ -chart, are computed from HARTMANN's formulas, reduced to 1945.0. They are:

$$\begin{aligned} X_n &= 18120 - 470\Delta\varphi + 45\Delta\lambda \\ Y_n &= -2315 + 30\Delta\varphi + 150\Delta\lambda \\ Z_n &= 43448 + 500\Delta\varphi \end{aligned}$$

with $\Delta\varphi = \varphi - 52^\circ$ and $\Delta\lambda = \lambda - 5^\circ$.

The differences $X - X_n$, $Y - Y_n$ and $Z - Z_n$ between the measured values X , Y and Z and the normal values X_n , Y_n and Z_n are used for drawing the chart of anomalies. The values $X - X_n$ and $Y - Y_n$ are compounded into a resultant horizontal disturbance vector, the values of $Z - Z_n$ given in γ are printed below the stations.

On the Belgian territory the isanomalics are copied from HARTMANN's chart, they have been computed from a normal equation valid for 1913.0 viz.

$$Z_n = 43156 + 530\Delta\varphi - 85\Delta\lambda$$

this formula being derived by a method analogous to that used for the formulas for the Netherlands.

¹⁾ A. D. LEWIS, Magnetic declination in South Africa (1936), Department of Irrigation, Pretoria.

²⁾ G. S. LJUNGDAHL, A magnetic survey of Sweden, Kungl. Sjökarteverket, Jordmagnetische Publikationer No. 9, Stockholm 1934.

³⁾ PH. C. P. HARTMANN, l.c.

For Germany a formula for the normal Z-field is given by RÖSSIGER⁴⁾. This reads:

$$\begin{aligned} Z_n &= 43460 + 499.9\Delta\varphi + 19.28\Delta\lambda + 0.952\Delta\varphi\Delta\lambda + \\ &\quad - 10.95(\Delta\varphi)^2 + 2.21(\Delta\lambda)^2 \end{aligned}$$

with $\Delta\varphi = \varphi - 52^\circ.07$, $\Delta\lambda = \lambda - 12^\circ.68$ valid, for 1941.5. This normal field is constructed from the yearly means at the magnetic observatories Niemegek, San Fernando, Eskdalemuir, Lovö, Kasan and Helwan and is, therefore, applicable for the whole of Europe. The annual changes of the coefficients are: for the term with $\Delta\varphi$: -0.60 , for the term with $\Delta\lambda$: $+1.61$, for the term with $\Delta\varphi\Delta\lambda$: -0.058 , for the term with $(\Delta\varphi)^2$: $+0.42$, for the term with $(\Delta\lambda)^2$: -0.206 . Neglecting the terms of higher order, and reducing this formula to $\varphi = 52^\circ$, $\lambda = 5^\circ$ as a magnetic centre, the approximation for the epoch 1945.0 becomes:

$$Z_n = 43470 + 499\Delta\varphi + 21\Delta\lambda$$

This formula differs only slightly from the one used in this publication, apart from the term with $\Delta\lambda$, which is of little importance.

For the northeastern borderland REICH's⁵⁾ charts of the anomalies of the vertical component were available. Moreover, the isanomalics of the disturbed region of Lingen have already been published by REICH⁶⁾. These isanomalics show a close correspondence with ours and have been drawn in the northeastern borderland of the Δ -map. For the southeastern region no German maps were available.

CHAPTER IV

DISCUSSION OF THE CHARTS

The D-chart.

The isogonics of the declination chart have been drawn at mutual distances of 5'. As the uncertainty is greater on Belgian territory than on Dutch, they have there been drawn for every 10'. It was not possible to continue the isogonics around the province of South-Limburg. Over the North Sea the isogonics have been drawn in accordance with the general direction in NW-Europe; the accuracy of the lines is here much smaller than on land. We see that the declination varies from 6° 15' to 8° 0' between the east frontier of the Netherlands and the isle of Walcheren. A disturbed region extends from South-Limburg to Utrecht. The disturbances of the Belgian region turn out to be continued in the province of Zeeland and the west part of North Brabant. Some other important disturbances are found in the neighbourhood of the Dollart and in the Wadden Sea.

The H-chart.

The horizontal component varies from 17600 γ over the Wadden islands in the north to 18800 γ over South-Limburg. The disturbed region over the Wadden Sea and in West-Friesland is very striking. The disturbance over the Dollart becomes evident by an abnormally steep gradient of the horizontal component. The central part of the Netherlands

⁴⁾ M. RÖSSIGER, Beiträge der angewandten Geophysik, 9, 121, 1941.

⁵⁾ Thanks are due to Prof Dr. H. REICH for furnishing the maps.

⁶⁾ H. REICH, Zeitschr. der deutschen geologischen Gesellschaft 85, 935, 1933.

is only slightly disturbed. The Belgian region together with South-Limburg and Zeeland show a strong irregularity as is also the case on the *D*-chart.

The *Z*-chart.

The vertical component increases between South-Limburg and the Wadden islands from 42850 γ to 44200 γ . The chart shows remarkable disturbances in the Wadden Sea and in the area of the Dollart. The disturbed region from South-Limburg to Utrecht is characterized by a downward bulging of the isolines, which means an abnormally large value of the vertical component or a positive anomaly. Just as in the case of the *D*-chart and the *H*-chart the south part of the *Z*-chart is strongly disturbed.

The *X*-chart.

The chart of the northerly component is naturally very much like the *H*-chart and, therefore, does not give rise to special remarks. The accuracy of the *X*-chart is determined by the formula

$$X = H \cos D, \text{ from which: } dX = \cos D \cdot dH - H \sin D \cdot dD$$

Over the whole country this formula can be simplified to

$$dX = dH - 0,5 dD$$

(*dX* and *dH* in γ , *dD* in minutes of arc) from which it follows that the accuracies of *X* and of *H* are practically the same.

The *Y*-chart.

The isodynamics of the easterly component have a nearly north-south direction. In the neighbourhood of the east frontier the value is -1950 γ , over Walcheren it amounts to -3000 γ . The disturbed regions of the *Y*-chart are essentially the same as of the other charts. The accuracy of the *Y*-values follows from

$$Y = H \sin D, \text{ from which: } dY = \sin D \cdot dH + H \cos D \cdot dD$$

This can be reduced approximately to:

$$dY = 0,1 dH + 5 dD$$

(*dY* and *dH* are expressed in γ , *dD* in minutes of arc). The accuracy of *Y* is in the first place determined by that of *D*, an error of 1' has an influence of 5 γ upon *Y*.

The air-earth current.

In a publication "Notes on isomagnetic charts: VI Earth-air electric currents, and the mutual consistency of the *H* and *D* isomagnetic charts" (CHAPMAN¹⁾) concludes that a non-potential part of the geomagnetic field must be attributed to errors in the values of *D* and *H*. This means, that the value of the vertical downward electric current *I* from the air into the earth, derived from the line integral of *H* along a closed curve

$$\int H ds = 4\pi I$$

must be very small or even zero. In order to investigate this three trapezia 1, 2 and 3 of 5435, 5100 and 8325 km² resp. have been drawn with their sides parallel to the circles of longitude and latitude. Along the parallel sides of these trapezia the mean value of *Y* was determined over intervals of 5 km and also the value of *X* along the slanting sides.

¹⁾ S. CHAPMAN, Notes on isomagnetic charts, Terr. Mag. 47, 1, 1942.

The integral of *H* along these trapezia, taken in clockwise direction, gives for 1 the value + 1045 oersted cm, for 2 : -1550 oersted cm and for 3: +1415 oersted cm. From these values the following currents are calculated: in 1: +0.15 A/km², in 2: -0.24 A/km², in 3: +0.14 A/km² (+ means a downward current).

After SCHMIDT's map²⁾ the intensity of the air-earth current is -0.2 A/km² in NW-Europe. CHAPMAN's³⁾ values differ considerably from SCHMIDT's map.

The difference between the directions of the currents in the adjoining regions 1 and 2 raises the surmise that these currents are not real. In order to investigate this in more detail each trapezium is subdivided into four parts and in each part the strength of the current is calculated separately. The result is:

$$1a: -0.43 \text{ A/km}^2, 1b: +0.02 \text{ A/km}^2, 1c: +0.48 \text{ A/km}^2, 1d: +0.64 \text{ A/km}^2.$$

$$\text{In } 2a: -0.73 \text{ A/km}^2, 2b: +1.75 \text{ A/km}^2, 2c: -0.84 \text{ A/km}^2, 2d: -0.90 \text{ A/km}^2.$$

$$\text{In } 3a: +3.78 \text{ A/km}^2, 3b: -2.56 \text{ A/km}^2, 3c: -0.23 \text{ A/km}^2, 3d: -0.22 \text{ A/km}^2.$$

It turns out that the currents per km² become stronger according as the integrals enclose a smaller area, and that there is not any regularity in the direction of the currents. This points evidently to the influence of the accidental errors in *X* and *Y*, which get more important according as a smaller region is considered.

We will conclude with the remark that a periodically variable air-earth current cannot be detected by the above mentioned method. The measurements have been carried out on various days and times, and have been reduced without taking into consideration an eventual local current. This local current could be found from simultaneous measurements along a closed line. The only statement that can be made is that a permanent air-earth current does not exist.

The Δ -chart.

The deviations from the normal field, ΔX , ΔY and ΔZ , have been collected in one chart. The horizontal anomalies ΔX and ΔY have been composed to form one vector, the value ΔZ is printed in γ . The isanomals of ΔZ are drawn for every 20 γ . The regions with positive values of ΔZ are dotted red, the negative regions are white. Only a few lines have been drawn over Belgian territory, but in part of the German borderland we were able to complete the chart.

The *G*-chart.

The last chart is the chart of the gravity anomalies in the Netherlands, based on measurements of the Government's Commission for Geodesy, the Bataafsche Petroleum Maatschappij and the Government's Mines. This chart was kindly put at our disposal by the Bataafsche Petroleum Maatschappij. The isogams have been drawn for every 5 mgal. The regions with an abnormally large value of the gravity are dotted red, the negative regions are white.

²⁾ J. BARTELS and S. CHAPMAN, Geomagnetism, page 665.

³⁾ S. CHAPMAN, l. c.

The magnetic and gravimetric anomalies.

The magnetic anomalies are caused by irregular deviations from a horizontally homogeneous magnetization in the earth's crust. Our knowledge of the details and quantitative values for magnetization of rocks is still in a rather unsatisfactory state owing to the lack of experimental determinations of the magnetic properties of rocks in low magnetic fields. A factor of primary importance is the magnetic susceptibility, but this factor depends highly on the strength of the magnetic field in which it is measured. Many values given in the literature refer to a field-strength larger than that of the earth's magnetism.

A great many values of the susceptibility of various rocks are given by HAALCK ¹⁾, HEILAND ²⁾ and NETTLETON ³⁾. It is probable that the magnetic properties of most rocks are determined by magnetite (Fe₃O₄) which is present almost everywhere in the form of an admixture.

The susceptibility and also the residual magnetism of the crystalline rocks are much larger than the corresponding values for sedimentary rocks. The ratio of the susceptibilities is of the order of 100 to 1000. This convinced REICH ⁴⁾ that at least in North Germany the archaic crystalline massifs and the crystalline rocks of the Palaeozoicum cause the extensive positive regional disturbances of the magnetic field, whereas the sedimentary rocks are accompanied by negative anomalies, because the small value of the susceptibility of the sediments is compensated by the greater depths of the crystalline rocks. NETTLETON concludes that the measurements may be considered as being made on a plane (i.e. the surface of the ground) which is supported on a nonmagnetic medium (i.e. the sediments) and that

the effects measured are caused by undulations of the magnetic surface (i.e. the surface of the igneous rocks) and polarization-contrasts below that surface. Generally speaking, the magnetic anomalies in sedimentary regions give an image of the tectonics of the crystalline substructure. Where the basement rock in the form of a horst or a massif is higher than in its neighbourhood, it will usually cause a positive disturbance. The magnitude of this disturbance depends on the susceptibility of the rock and on its depth.

The interpretation of magnetic anomalies is not unique. As not only the susceptibilities of the rocks but also their dimensions and depths exert an influence, a given anomaly can be explained by a variety of causes. The right interpretation can, therefore, only be obtained from additional dates (geological, gravimetric and seismic).

In fig. 11 taken from NETTLETON (page 221), the calculated profile has been drawn of a magnetic anomaly, caused by a long ridge with vertical sides, representing an uplift of one-fourth of the depth and with a width equal to the thickness of the sediment. The polarizing field is taken vertical and the susceptibilities of sediment and basement are supposed to be $k = 0$ and $k = 6 \cdot 10^{-3}$ respectively, so that the magnetic polarization I

$$I = kZ$$

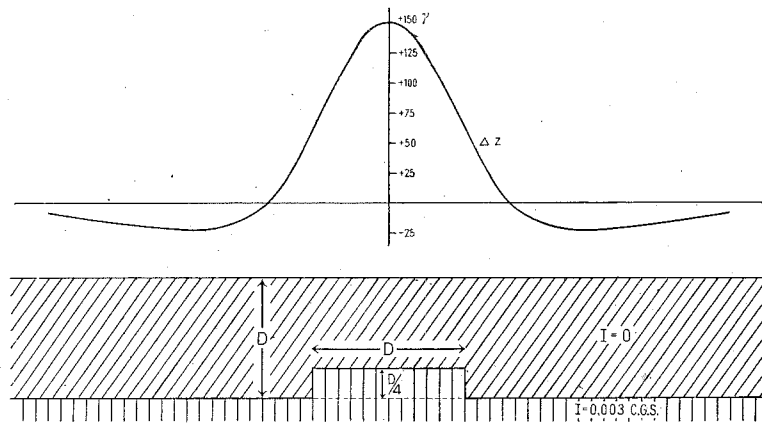


Fig. 11. Vertical magnetic anomaly caused by a long ridge (after Nettleton.)

with a field strength $Z = 0.5$ oersted is equal to 0 for the sediment and to 0.003 c.g.s. for the basement. The calculation of the magnetic influence of the horst is carried out in the simplest way by replacing the volume-magnetization by a uniform distribution of magnetic poles with a density of I per cm² along the upper and lower sides of the horst. It turns out that the positive disturbance over the horst with a value of 150γ is accompanied on both sides by weak negative disturbances, so that the value integrated over an infinitely extended horizontal surface is zero.

The geological interpretation of the magnetic anomalies

is given for the greater part by HARTMANN ⁵⁾ in his doctoral thesis. We borrow from it: "The old crystalline massifs of Brabant, Stavelot, Hohe Venn and Rocroy stand out with positive anomalies against the negatively disturbed surroundings. A zone with positive maxima extends from the Yser to Maastricht. The upward curving of the Axe d'Ostende is evident from the northward deflection of the isanomals. The nose of the massif of Brabant causes a positive anomaly in ΔZ of 225γ near Maastricht in accordance with the small depth at which the Cambro-Silur is here found

⁵⁾ PH. C. P. HARTMANN, Thesis.

directly under the Senon. The "Sillon du Rupel-Demer" is marked by a strip of negative anomalies from Hasselt to the West. North of this strip we find a positive maximum near Antwerp. The massif of Stavelot-Hohe Venn causes a strong positive disturbance. To the north of this massif the crystalline rock is covered by devon, carbon and chalk. We see therefore a strong decrease of the anomaly to the north so that even a negative value of -14γ is found near Eysden. The influence of the massif of Brabant is furthermore expressed on Dutch territory in the strong positive anomaly of Woensdrecht".

The measurements carried out later on in Zeeland show that the greater part of Zeeland is strongly disturbed by positive anomalies, from which we infer that the limit of the massif of Brabant presumably follows a line from the South-Holland islands to South-Limburg. In this region all disturbance vectors of the horizontal field are directed perpendicularly to this line.

About the magnetic anomalies on Dutch territory HARTMANN tells us: "Most striking is the Netherlands direction which dominates the tectonics of our country. Regions of positive and negative anomalies alternate and show very clearly the NW-SE direction. The Swabian direction can also clearly be seen on the chart, namely from the shifting of the isanomals along WSW-ENE directed lines.

The Central Graben can be clearly seen as a negatively disturbed region. We find the minimum in the province of North Brabant south of Eindhoven. More to the NW the zone, which is magnetically negative, seems to shift to the NE. The Peelhorst and its continuation, the horst of Erkelenz, cause a positive anomaly. To the north this anomaly is bordered by a negative disturbance, by which the troughs of Venlo and Venray are represented. The horst of Viersen deflects the isanomal of $+20\gamma$ to the NW again, where upon it curves round the horst of Mill, which is shifted to the NE with regard to the Peelhorst even in its magnetic anomaly."

Comparing now the Δ -chart with the G -chart we see that the gravimetric chart too gives a clear image of the Central Graben; the gravimetric minimum, however, is situated more to the north than the magnetic one. The strip of positive magnetic anomalies which borders the Central Graben to the east, coincides with the region of horsts, as published in the *Mijnbouwkundig Jaarboek 1947*¹⁾, but is found at the limit of the gravimetric negative region. The gravimetrically deep trough of the Roer can be found on the magnetic chart too as a shallow between the maxima, but does not lead to negative values. As regards the region west of the Central Graben the two charts agree well with each other.

HARTMANN continues: "One gets the impression from the magnetic anomalies that a region of grabens extending over the Rhine is situated between the horst of Mill with its continuation to the north, and the high region of the Achterhoek. A partly broad region of horsts is sharply outlined by the magnetic anomalies in the continuation of the zone Erkelenz—Peel—Mill over "het Gooi", the

Utrecht ridge, an important part of the Gelderse Vallei and even of the Neder Veluwe to Voorthuizen. Over the horst of the Utrecht ridge the disturbance values are highest and this horst seems to extend farthest to the NW and the SE. The geomagnetic anomalies in this region support the hypothesis that the existence of the Utrecht ridge is not primarily due to the pushing of the land-ice, but that a tectonically caused higher structure accounts for the formation of a push-moraine. The horst of the Utrecht ridge is shifted to the NE relative to the horst of Mill. The same was found for the situation of the horst of Mill in relation to the Peelhorst, where the shift took place along the fault-lines of the transverse graben of Venray. It is therefore possible to suppose the existence of a transverse graben following the line Tiel—Dieren. The sudden deflection of the rivers is perhaps partly caused by this graben."

We can remark in this connection that this transverse graben is still better visible in the Δ -chart than in HARTMANN's chart. The magnetic maximum near the Utrecht ridge is not visible on the G -chart, the spur to the Neder-Veluwe, however, coincides with a slight gravimetric maximum of $+5$ mgal.

HARTMANN continues:

"To the east of the magnetically positive disturbed region of the Utrecht ridge lies a negative region. Probably the Rhine Graben which could be followed to the Ysel east of Arnhem is continued in the NW direction, slightly shifted to the NE. The east border of this graben may be supposed to lie near Apeldoorn. East of it the high ground of the east part of the Netherlands begin probably to rise stepwise. The west border of the graben is supposed to follow a line drawn over Voorthuizen".

The graben, supposed to exist by HARTMANN, is very clear from a gravimetric viewpoint. Besides, we see that on the G -chart a minimum region in the province of North-Holland is visible as a continuation of this Rhine graben. Also from a magnetic viewpoint this continuation in NW-direction through the centre of North-Holland is quite evident. HARTMANN remarks about the magnetic anomalies in the province of South-Holland:

"Parallel to the zone with positive magnetic anomalies in "het Gooi" and Utrecht, a zone with negative values runs SW of it through Nieuw Vennepe and Vreeswijk. Further to the SW the anomalies increase first, but soon they decrease again so that a second zone of minima seems to run from Leidschendam over Zevenhuizen in the direction of Gorinchem. It is obvious to suppose a secondary horst to lie between these two more negatively disturbed regions".

Further, according to HARTMANN a graben can faintly be perceived near the Nieuwe Waterweg and he attempts to connect it with the bed of the original Rhine, as constructed from the location of fossils of *Viviparus glacialis*. This graben is not visible on the complete Δ -chart; however, we see a strongly curved minimum region over the South-Holland islands, bordered in the north by a small maximum over the Waterweg. The corresponding gravimetrically negative region is situated more northward and its axis coincides with Nieuwe Waterweg. Further we see on the G -chart a maximum extending from The Hague

¹⁾ W. J. VAN RIEL, Geophysische metingen ten behoeve van de Kolennijbouw, *Mijnbouwkundig Jaarboek 1947*.

in SE-direction to the Lek. This sustains HARTMANN's supposition of a secondary horst in the graben: the complete Δ -chart shows a positive strip between the Hague and Haarlem.

HARTMANN remarks about the province of North-Holland: "The whole of the province is negatively disturbed in a magnetic sense. Nearly the whole province shows negative anomalies of the gravity as well. We find the minimum of the magnetic anomaly near Alkmaar and Hoorn. FABER's geological profile which indicates a broad graben, is in accordance with this minimum. This whole province belonged to the small region, which was flooded by the Eemsea. On the other hand the sea of the High-terrace did not penetrate here, but did so into the graben south of the town of The Hague, which remained dry during the irruption of the Eemsea.

To the northeast the value of the anomaly of the vertical component increases again and near Staveren (station 74) a positive anomaly of $+15\gamma$ is found. The basement rock probably rises stepwise in that direction.

The NW-SE line, drawn SW along the isles of Texel and Wieringen, formed the shore of the sea of the High-terrace, the Eemsea and the sea of the Boreal-Atlanticum. In the first case the isles of Texel and Wieringen were flooded, in the two other cases they remained dry, whereas the sea flooded the province of North-Holland (TESCH 1939).

It is obvious, therefore, to suppose the existence of a faultline somewhere near here, which borders the Alkmaar graben to the northeast. The diluvium is situated higher in the isles of Texel and Wieringen than elsewhere in the province of North-Holland. The tectonically higher structure of Texel may at the same time explain the fact that the coast-line juts out in a northwest direction, a phenomenon which appeared on a larger scale in the situation of the original reefs. Probably Gaasterland forms also part of this zone of horsts."

Let us now have a look at the centre of East Netherland. Beyond the German frontier we see a strong magnetical and gravimetrical maximum, the maximum of Lingen, which makes its influence felt within the Netherlands. This is part of the crystalline massif of Bramsche, situated more towards the east. The high grounds of the Achterhoek are magnetically as well as gravimetrically disturbed by positive variations. The magnetic minimum in the centre of the Achterhoek nearly coincides with a saddle-region between two positive regions in the G -chart. It is obvious therefore, to suppose here the existence of a graben, as indicated already by VAN WATERSCHOOT VAN DER GRACHT between Buurse and Rekken. Northwest of it a magnetically negative strip extends through West-Overijssel and South-west Friesland, in which all isanomals are stretched out in a northwest direction. The corresponding region in the gravimetric chart is positive and points to a maximum over South Friesland and the head of Overijssel. Here the gravimetric and the magnetic maps are contradictory. HARTMANN writes about this:

"Since the Jura the high grounds of the Achterhoek and Twente have sunk considerably less than the region north of it, where very thick mesozoic and tertiary layers were

deposited. If we assume that the gravimetric horst (here the maximum over the northwestern part of Overijssel is meant) is relatively young, its abnormal magnetic field can be explained by the great thickness of the calcareous mesozoic sediments. The susceptibility of limestone and marl is very small, much smaller than the susceptibility of other sediments. Their density, however, is greater than the density of sands and clays. Owing to the rising of the horst these very thick mesozoic layers have come to lie higher than in their surroundings. By their greater density they cause a positive gravity anomaly. However, by their smaller susceptibility these layers can cause a negative magnetic disturbance. The crystalline basement, which is situated at a rather great depth notwithstanding the rising, has a smaller influence upon the magnetic field, so that a negative anomaly may be the result of the elevation".

It is also possible to conceive the presence of a graben in the crystalline basement, filled with sediments of great density, which are responsible for the gravity maximum. This interpretation is simpler than HARTMANN's. A decision between the two is not possible for the time being, through lack of knowledge of the deeper substructure. A negative strip runs through the provinces of Drenthe and Groningen in the Netherlands direction SE-NW, which it is possible to follow in the Wadden Sea north of Groningen. The G -chart too shows a negative region getting deeper in the NW direction. At the northeast side this strip is bordered by an extensive magnetic maximum, the oblong centre of which is situated over the Eems and the Dollart. Presumably a crystalline ridge can be found here, which follows the graben through the province of Groningen. HARTMANN remarks about this graben that the Eemsea-transgression has been able to penetrate here only along a narrow bay from the Lauwersea to the Lake of Zuidlaren. The region west of this graben, the greater part of the province of Friesland and the Wadden Sea show a positive magnetic anomaly. The gravity anomaly is negative here though an increase of the gravity must exist from NW-Friesland to the isles of Vlieland and Terschelling. Just as in South-Friesland a contradiction is also present in North-Friesland between the magnetic and the gravimetric maps. Most remarkable is the narrow tongue of positive magnetic anomalies penetrating from the Wadden Sea in an eastward direction into Friesland. A narrow ridge in the crystalline basement may be responsible for it. The isolated negative anomaly on the isle of Ameland is not quite certain, it may be caused by an error.

Finally we see in the Wadden Sea between the isles of Texel, Vlieland, Terschelling and the Friesian coast a very strong and steep magnetic maximum, the anomaly in the centre situated at the Inschot fairway reaching a value of more than 200γ . The cause of this strong disturbance, visible in all magnetic charts, is completely unknown. Up to now neither gravimetric nor seismic investigations have been carried out in this region, but it will be very interesting to compare the results of these investigations, which we hope will not be postponed much longer, with the magnetic anomalies. The limited area of this disturbance, the steep gradient and the high value of its maximum are only comparable with the disturbances near the south

frontier of our country, where the crystalline rocks with a considerable magnetite content are found close under the surface.

Magnetic anomalies and seismicity.

By J. P. ROTHÉ ¹⁾ attention is drawn to the fact, that in some cases a relation can be found between the epicenters of earthquakes and magnetic anomalies. It was found in Japan that the epicenters of some recent heavy earthquakes were situated in magnetically strongly disturbed regions, where moreover the geomagnetism showed an abnormally strong secular change. In France, too, centres of earthquakes are found on the axis of the magnetically Basin of Paris and here also important changes of the anomalies occurred from the year 1896 to 1924. ROTHÉ ascribes this fact to magmatic displacements which at the same time cause the earthquakes and the changes of the geomagnetic field.

In the Netherlands earthquakes are very rare phenomena. We borrow from the surveys of VISSER ²⁾ and VAN RUMMELEN ³⁾ the following. In the years 1833, 1850 and 1883 shocks have been felt in the provinces of North- and South-Holland with epicenters in the neighbourhood of Haarlem. Quakes are also mentioned near Harderwijk in the years 1781, 1824, 1859 and 1906. Furthermore in the year 1829 a shock was observed in Zwolle and in the year 1843 in Blokzijl. In the southern part of our country earthquakes are more numerous than in the northern part. Seve-

ral epicenters were found on the eastern border of the Central Graben: Tiel (1928, 1932), Uden and Veghel (1843, 1848, 1932), Dinter (1848), Vught (1932), Asten (1932), Gratem (1906), Roermond (1851). The town of Sittard situated on the western border of the Central Graben experienced a very slight earthquake in the years 1918 and 1931. Several times shocks have been observed related with movements along the Feldbiss, an important tectonical disturbance in the region of the Limburg mines: Herzogenrath (1873, 1874, 1877) and Rolduc (1928). Also related to faults are the quakes of Heerlen (1914) and Voerendaal (1930). It is evident therefore that even in recent years movements occur along the fault planes in the South of the Netherlands. The most important shocks are connected with shifts in the fault-region between the Central Graben and the Peelhorst. The epicentres, however, are not found in the axis of the magnetic anomaly as in France, but along the western border of the related magnetic maximum. To what extent sudden changes in the magnetic anomalies are caused by the tectonical movements is a question which cannot be answered as the surveys of 1892 and 1945 differ too much in accuracy. Probably, the shocks in the northern part of the Netherlands are also of tectonical origin. The epicenters of Harderwijk, Zwolle and Blokzijl are indeed situated in transition-regions between magnetic maxima and magnetic minima and are perhaps related to movements along the borders of the Alkmaar graben. The Δ -chart and the G -chart give the impression that this graben extends via the Veluwe through the province of North-Holland as a continuation of the Rhine graben, which from a seismic viewpoint is still active. It is not possible to relate the shocks in the neighbourhood of Haarlem to a gravimetric or magnetic structure.

¹⁾ J. P. ROTHÉ, Tremblements de terre et anomalies magnétiques, *Geofisica pura et applicata*, 12, 1, 1948; Thèses, Paris, 1937.

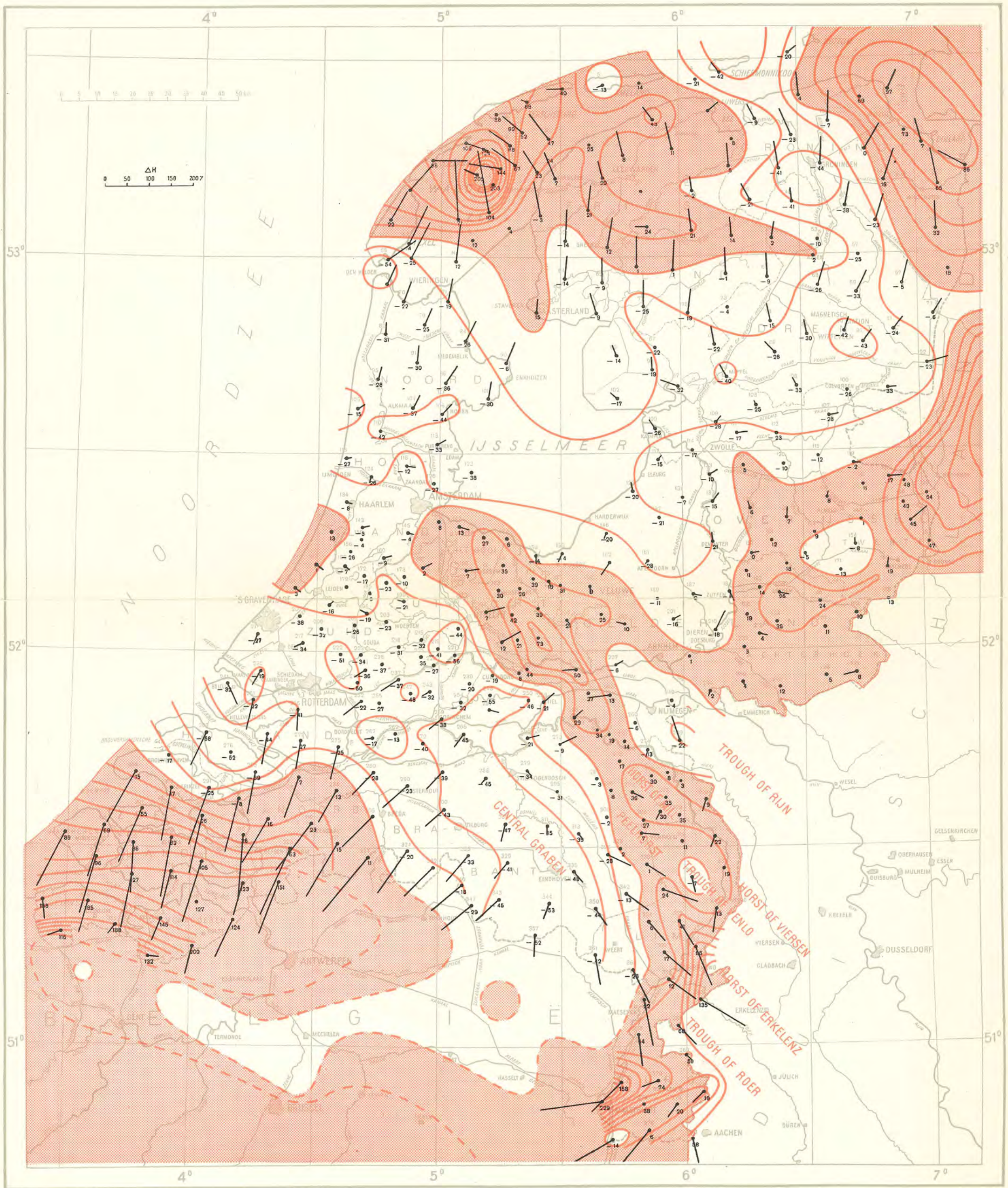
²⁾ S. W. VISSER, Aardbevingen in Nederland, *Tijdschrift Ned. Aardr. Genootschap*, 59, 494, 1942; *Seismologie*, Gorinchem, 1949.

³⁾ F. H. VAN RUMMELEN, Overzicht van aardbevingen enz. *Mededelingen Jaarverslag Geologisch Bureau te Heerlen*, 1945.

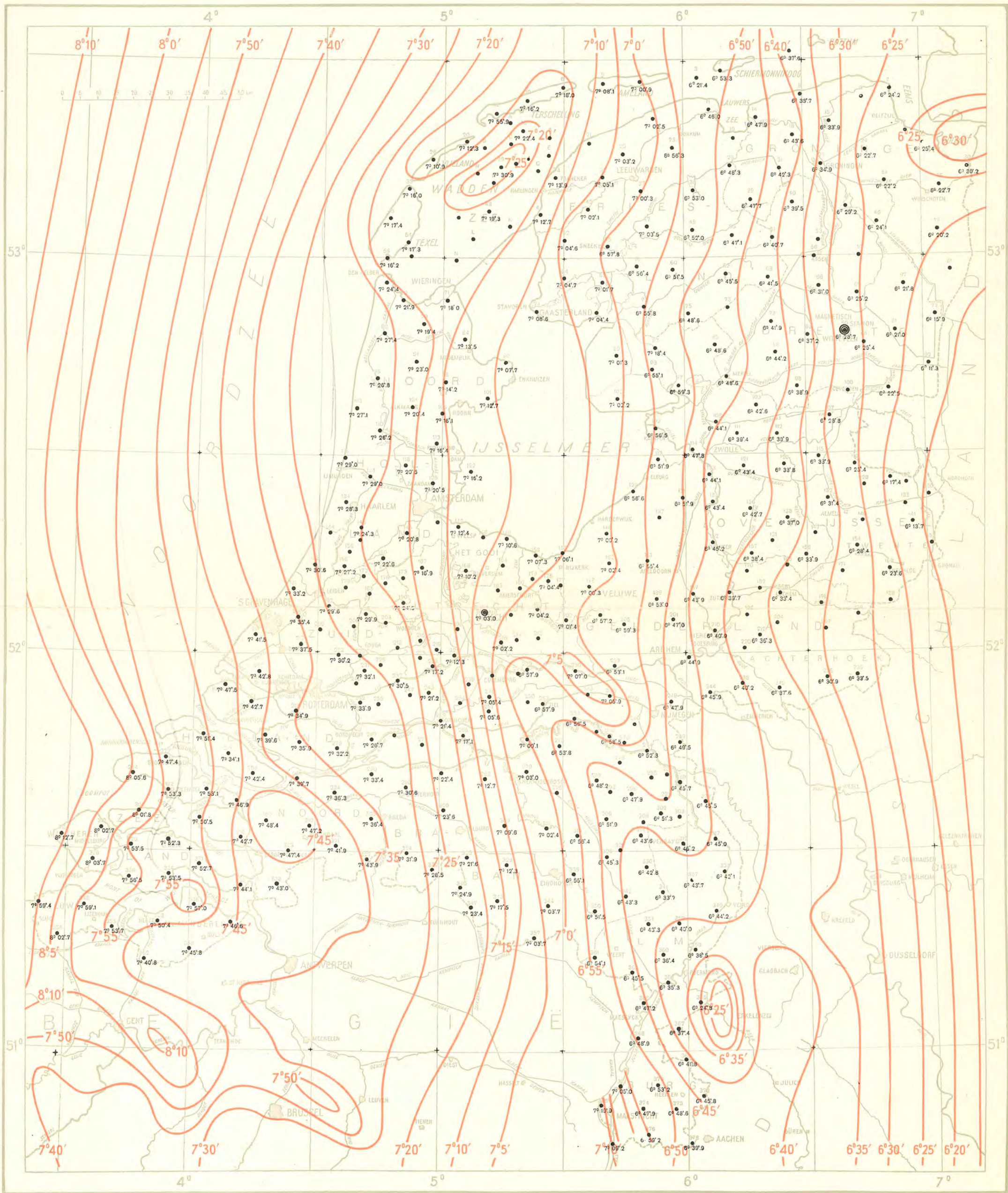
	Name of station	φ	λ	-D	H	X	-Y	Z	ΔX	ΔY	ΔZ
84	Medemblik	52° 47'.70	5° 05'.32	7° 13'.5	17937γ	17794γ	2256γ	43819γ	+ 44γ	+ 22γ	- 26γ
85	Steenwijk	52° 46'.90	6° 07'.79	6° 48'.6	17969γ	17843γ	2131γ	43817γ	+ 40γ	- 9γ	- 22γ
86	Benneveld	52° 46'.88	6° 44'.51	6° 25'.4	17971γ	17858γ	2010γ	43796γ	+ 28γ	+ 21γ	- 43γ
87	Blankenham	52° 46'.45	5° 52'.93	7° 18'.4	17969γ	17823γ	2285γ	43813γ	+ 27γ	- 125γ	- 22γ
88	Ruinen	52° 44'.48	6° 22'.71	6° 44'.2	17980γ	17856γ	2109γ	43793γ	+ 23γ	- 23γ	- 26γ
89	Rutten	52° 45'.18	5° 43'.12	7° 01'.3	17956γ	17822γ	2195γ	43810γ	+ 24γ	- 11γ	- 14γ
90	Andijk	52° 44'.64	5° 15'.67	7° 07'.7	17973γ	17834γ	2230γ	43814γ	+ 52γ	+ 24γ	- 6γ
91	Nieuwe Niedorp	52° 44'.20	4° 53'.08	7° 23'.0	17954γ	17805γ	2307γ	43786γ	+ 36γ	+ 3γ	- 30γ
92	Klazienaveen	52° 43'.55	7° 00'.63	6° 11'.3	17986γ	17881γ	1939γ	43788γ	+ 12γ	+ 53γ	- 23γ
93	Marknesse	52° 42'.95	5° 52'.05	6° 55'.1	17983γ	17852γ	2166γ	43787γ	+ 30γ	- 2γ	- 19γ
94	Meppel	52° 41'.92	6° 10'.53	6° 48'.6	17999γ	17872γ	2134γ	43758γ	+ 28γ	+ 16γ	- 40γ
95	Koedijk	52° 41'.70	4° 43'.73	7° 26'.8	17958γ	17807γ	2327γ	43767γ	+ 26γ	+ 8γ	- 28γ
96	Wognum	52° 41'.02	5° 00'.47	7° 14'.2	17980γ	17837γ	2265γ	43754γ	+ 38γ	+ 28γ	- 36γ
97	Vollenhove	52° 40'.50	5° 58'.39	6° 59'.3	17993γ	17864γ	2190γ	43753γ	+ 17γ	- 41γ	- 32γ
98	Kerkenbosch	52° 40'.47	6° 28'.03	6° 38'.9	18010γ	17889γ	2085γ	43752γ	+ 20γ	- 10γ	- 33γ
99	Oud-Schoonebeek	52° 39'.82	6° 50'.64	6° 22'.5	18004γ	17893γ	1999γ	43747γ	+ 2γ	+ 20γ	- 33γ
100	Coevorden	52° 39'.67	6° 40'.59		18008γ			43753γ			- 26γ
101	Oosterleek	52° 38'.66	5° 10'.89	7° 12'.7	18008γ	17866γ	2261γ	43740γ	+ 41γ	+ 8γ	- 30γ
102	Nagele	52° 38'.49	5° 43'.33	7° 02'.2	18001γ	17866γ	2205γ	43752γ	+ 15γ	- 17γ	- 17γ
103	Oud-Avereest	52° 37'.62	6° 17'.82	6° 42'.6	18017γ	17894γ	2105γ	43736γ	+ 11γ	- 4γ	- 25γ
104	Ursem	52° 37'.31	4° 52'.23	7° 20'.4	17995γ	17847γ	2299γ	43722γ	+ 25γ	+ 16γ	- 37γ
105	Egmond aan Zee	52° 37'.17	4° 38'.47	7° 27'.1	17978γ	17826γ	2332γ	43743γ	+ 13γ	+ 18γ	- 15γ
106	Schardam	52° 36'.20	4° 59'.81	7° 16'.1	18003γ	17858γ	2278γ	43706γ	+ 22γ	+ 19γ	- 44γ
107	Collendoorn	52° 35'.74	6° 36'.04	6° 28'.8	18030γ	17914γ	2035γ	43718γ	+ 2γ	+ 22γ	- 28γ
108	Hasselt	52° 35'.14	6° 07'.82	6° 44'.1	18036γ	17911γ	2115γ	43713γ	+ 15γ	+ 13γ	- 28γ
109	Kampen	52° 34'.41	5° 52'.76	6° 56'.5	18042γ	17910γ	2181γ	43709γ	+ 20γ	- 15γ	- 26γ
110	Akersloot	52° 33'.76	4° 44'.33	7° 26'.2	18002γ	17850γ	2330γ	43687γ	+ 7γ	+ 7γ	- 42γ
111	Nieuw-Leuzen	52° 33'.22	6° 13'.10	6° 39'.4	18036γ	17914γ	2091γ	43708γ	0γ	+ 24γ	- 17γ
112	Ommen	52° 33'.20	6° 22'.72	6° 38'.9	18045γ	17924γ	2089γ	43702γ	+ 2γ	+ 2γ	- 23γ
113	Kwadijk	52° 31'.75	4° 58'.23	7° 16'.4	18025γ	17880γ	2282γ	43680γ	+ 10γ	+ 21γ	- 33γ
114	Zwolle	52° 30'.71	6° 01'.98	6° 47'.8	18059γ	17932γ	2137γ	43686γ	+ 7γ	+ 8γ	- 17γ
115	Mariënberg	52° 29'.61	6° 33'.40	6° 33'.9	18067γ	17949γ	2066γ	43683γ	- 9γ	+ 1γ	- 12γ
116	Wijk aan Zee	52° 29'.55	4° 35'.90	7° 29'.0	18025γ	17872γ	2347γ	43667γ	+ 2γ	+ 13γ	- 27γ
117	Oosterwolde	52° 29'.22	5° 53'.65	6° 51'.9	18070γ	17940γ	2160γ	43676γ	+ 9γ	+ 6γ	- 15γ
118	Zaandijk	52° 28'.55	4° 50'.67	7° 20'.5	18033γ	17885γ	2304γ	43674γ	- 4γ	+ 20γ	- 12γ
119	Bruinehaar	52° 28'.49	6° 42'.23	6° 25'.4	18086γ	17972γ	2023γ	43683γ	- 2γ	+ 22γ	- 2γ
120	Lemelerberg	52° 28'.39	6° 24'.48	6° 38'.8	18077γ	17956γ	2092γ	43675γ	- 5γ	- 2γ	- 10γ
121	Hoonhorst	52° 28'.23	6° 14'.70	6° 43'.4	18079γ	17955γ	2117γ	43688γ	0γ	- 3γ	+ 5γ
122	Marken	52° 27'.68	5° 06'.20	7° 16'.2	18051γ	17906γ	2284γ	43641γ	- 2γ	+ 1γ	- 38γ
123	Wapenvelde	52° 26'.92	6° 06'.08	6° 44'.1	18098γ	17973γ	2122γ	43662γ	+ 14γ	+ 15γ	- 10γ
124	Beverwijk	52° 26'.73	4° 42'.02	7° 29'.0	18035γ	17882γ	2349γ	43645γ	- 16γ	- 2γ	- 26γ
125	Vasse	52° 26'.18	6° 50'.91	6° 17'.4	18109γ	18000γ	1984γ	43683γ	+ 2γ	+ 41γ	+ 17γ
126	Broek in Waterland	52° 25'.75	4° 57'.71	7° 20'.5	18055γ	17907γ	2307γ	43636γ	- 9γ	+ 1γ	- 27γ
127	Oud-Ootmarsum	52° 25'.55	6° 55'.01					43709γ			+ 48γ
128	Geesteren	52° 25'.23	6° 44'.51					43669γ			+ 11γ
129	Nunspeet	52° 24'.63	5° 46'.82	6° 56'.6	18110γ	17977γ	2189γ	43633γ	+ 15γ	- 3γ	- 20γ
130	Denekamp	52° 23'.56	7° 00'.13		18149γ			43708γ			+ 64γ
131	Heerde	52° 23'.46	5° 59'.54	6° 51'.9	18113γ	17983γ	2165γ	43636γ	+ 2γ	- 11γ	- 7γ
132	Wierden	52° 23'.20	6° 35'.47	6° 31'.4	18140γ	18023γ	2060γ	43649γ	+ 13γ	+ 4γ	+ 8γ
133	Veesen	52° 22'.85	6° 07'.07	6° 43'.4	18132γ	18008γ	2123γ	43623γ	+ 17γ	+ 13γ	- 15γ
134	Overveen	52° 22'.81	4° 36'.26	7° 28'.3	18074γ	17920γ	2350γ	43630γ	- 4γ	+ 13γ	- 8γ
135	Siemerink	52° 22'.26	6° 54'.71					43683γ			+ 49γ
136	Raalte	52° 21'.79	6° 15'.88	6° 42'.7	18146γ	18022γ	2121γ	43635γ	+ 16γ	- 7γ	+ 6γ
137	Tepelberg	52° 20'.36	5° 53'.62		18136γ			43597γ			- 21γ
138	Holterberg	52° 20'.18	6° 25'.07	6° 37'.0	18166γ	18045γ	2093γ	43623γ	+ 19γ	- 1γ	+ 7γ
139	Waveren	52° 19'.82	4° 58'.96					43621γ			+ 8γ
140	Zenderen	52° 19'.71	6° 43'.63					43613γ			+ 1γ
141	Oldenzaal	52° 19'.40	6° 56'.30	6° 13'.7	18193γ	18086γ	1974γ	43655γ	+ 31γ	+ 40γ	+ 45γ
142	Muiden	52° 19'.21	5° 04'.10	7° 12'.4	18108γ	17964γ	2272γ	43621γ	- 8γ	+ 23γ	+ 13γ
143	Hoofddorp	52° 18'.92	4° 40'.35	7° 24'.3	18113γ	17962γ	2334γ	43603γ	+ 5γ	+ 21γ	- 3γ
144	Hillegom	52° 18'.26	4° 32'.52		18104γ			43613γ			+ 13γ
145	Amstelveen	52° 18'.22	4° 51'.38	7° 20'.8	18116γ	17967γ	2316γ	43596γ	- 4γ	+ 12γ	- 4γ
146	Ermelo	52° 18'.10	5° 40'.69	7° 00'.2	18129γ	17994γ	2210γ	43579γ	- 15γ	- 6γ	- 20γ
147	Rijssen	52° 17'.98	6° 31'.90					43607γ			+ 9γ
148	Naarden	52° 17'.48	5° 10'.30					43621γ			+ 27γ
149	Blaricum	52° 17'.31	5° 16'.02	7° 10'.6	18130γ	17988γ	2265γ	43598γ	- 8γ	+ 1γ	+ 6γ
150	Nieuw Vennepe I	52° 17'.18	4° 39'.74					43587γ			- 4γ
151	Bofink	52° 16'.68	6° 21'.27					43599γ			+ 12γ
152	Deventer	52° 16'.60	6° 06'.75	6° 45'.2	18177γ	18051γ	2138γ	43565γ	+ 11γ	+ 2γ	- 21γ
153	Lossen	52° 15'.98	7° 00'.86					43628γ			+ 47γ
154	Delden	52° 15'.97	6° 42'.30	6° 28'.4	18197γ	18081γ	2052γ	43573γ	+ 9γ	- 1γ	- 8γ
155	Nijkerk	52° 15'.62	5° 29'.75	7° 06'.1	18133γ	17994γ	2242γ	43574γ	- 25γ	- 9γ	- 4γ
156	Nieuw Vennepe II	52° 15'.33	4° 37'.07					43550γ			- 26γ
157	Bathmen	52° 15'.03	6° 16'.03	6° 38'.4	18183γ	18061γ	2102γ	43573γ	+ 2γ	+ 15γ	0γ
158	Markelo	52° 14'.78	6° 29'.99	6° 33'.9	18182γ	18062γ	2079γ	43566γ	- 9γ	+ 4γ	- 5γ
159	Bunshoten	52° 14'.88	5° 23'.49	7° 07'.3	18140γ	18000γ	2249γ	43568γ	- 21γ	0γ	- 4γ
160	Kudelstaart	52° 14'.38	5° 45'.38	7° 22'.6	18150γ	18000γ	2330γ	43559γ	+ 4γ	+ 15γ	- 9γ
161	Apeldoorn	52° 13'.97	5° 50'.73	6° 55'.4	18170γ	18037γ	2190γ	43536γ	- 11γ	- 9γ	- 28γ
162	Garderen	52° 13'.55	5° 41'.39	7° 02'.4	18152γ	18016γ	2225γ		- 30γ	- 20γ	
163	Voorhout	52° 13'.40	4° 28'.38	7° 30'.6	18141γ	17986γ	2371γ	43561γ	- 5γ	+ 16γ	+ 1γ
164	Baarn	52° 13'.28	5° 15'.21					43594γ			+ 35γ
165	Verwolde	52° 13'.06	6° 25'.00					43575γ			+ 18γ
166	Abbenes	52° 13'.04	4° 36'.08	7° 27'.2	18143γ	17990γ	2353γ	43549γ	- 10γ	+ 15γ	- 7γ
167	Vinkeveen	52° 12'.97	4° 55'.00	7° 16'.9	18172γ	18025γ	2303γ	43558γ	+ 11γ	+ 18γ	+ 2γ
168	Loosdrecht	52° 12'.50	5° 05'.85	7° 10'.2	18172γ	18030γ	2268γ	43559γ	+ 4γ	+ 26γ	+ 7γ

	Name of station	φ	λ	-D	H	X	-Y	Z	ΔX	ΔY	ΔZ
169	Usselo	52° 12'.43	6° 50'.38	6° 23'.6	18240y	18126y	2031y	43570y	+ 20y	+ 2y	+ 19y
170	Eesterhold	52° 12'.29	6° 15'.27					43562y			+ 11y
171	Wegdam	52° 11'.98	6° 38'.70					43535y			- 13y
172	Rijnsaterwoude	52° 11'.63	4° 40'.68		18161y			43528y			- 17y
173	Mijdrecht	52° 11'.43	4° 50'.42		18169y			43533y			- 10y
174	Ham	52° 11'.02	5° 22'.27					43579y			+ 39y
175	Hoewelaken	52° 10'.98	5° 26'.16	7° 04'.4	18178y	18039y	2238y	43550y	- 15y	+ 7y	+ 10y
176	Zevenhoven	52° 10'.56	4° 46'.16		18178y			43513y			+ 23y
177	Klaarwater	52° 10'.22	5° 29'.29					43564y			+ 31y
178	Voorthuizen	52° 10'.14	5° 36'.34	7° 00'.3	18185y	18049y	2218y	43541y	- 19y	+ 1y	+ 9y
179	Hoogmade	52° 10'.05	4° 36'.25		18183y						
180	Bunt	52° 09'.74	5° 19'.20					43555y			+ 26y
181	Wassenaar	52° 09'.68	4° 23'.64	7° 33'.2	18167y	18010y	2388y	43532y	- 7y	+ 13y	+ 3y
182	Almen	52° 09'.57	6° 18'.12					43542y			+ 14y
183	Pijnenburg	52° 09'.51	5° 13'.94					43557y			+ 30y
184	Zutfen	52° 09'.06	6° 10'.75	6° 39'.7	18182y	18059y	2109y	43518y	- 43y	+ 24y	- 6y
185	Aarlanderveen	52° 08'.96	4° 42'.16		18193y			43514y			- 9y
186	Lochem	52° 08'.92	6° 23'.15	6° 33'.4	18226y	18107y	2081y	43548y	- 5y	+ 22y	+ 26y
187	Loenen	52° 08'.84	6° 01'.80	6° 43'.9	18219y	18093y	2136y	43519y	- 4y	+ 20y	- 2y
188	Buurse	52° 08'.57	6° 50'.66					43533y			+ 13y
189	Hoenderloo	52° 08'.11	5° 53'.13	6° 53'.0	18221y	18090y	2184y	43505y	- 6y	- 6y	- 11y
190	Zegveld	52° 07'.68	4° 50'.52	7° 24'.5	18208y	18056y	2348y	43491y	+ 3y	- 13y	- 21y
191	Borculo	52° 07'.33	6° 33'.40					43533y			+ 24y
192	Leiden	52° 07'.00	4° 32'.20	7° 29'.6	18207y	18052y	2374y	43490y	+ 8y	+ 7y	- 16y
193	Woudenberg	52° 06'.23	5° 23'.42	7° 04'.2	18211y	18072y	2241y	43539y	- 17y	+ 12y	+ 39y
194	De Bilt	52° 06'.06	5° 10'.62	7° 03'.0	18231y	18093y	2238y	43505y	+ 12y	+ 47y	+ 7y
195	Boskoop	52° 05'.87	4° 41'.20	7° 29'.9	18226y	18070y	2378y	43478y	+ 10y	- 19y	- 19y
196	Zeist	52° 05'.66	5° 16'.85					43537y			+ 42y
197	Lunteren	52° 05'.65	5° 39'.03	6° 57'.2	18240y	18106y	2208y	43520y	+ 1y	+ 7y	+ 25y
198	Wichmond	52° 05'.60	6° 15'.18					43514y			+ 19y
199	Rekken	52° 05'.51	6° 42'.46		18292y			43504y			+ 10y
200	Leidschendam	52° 05'.22	4° 24'.46	7° 35'.4	18196y	18036y	2403y	43454y	- 17y	- 2y	- 38y
201	Woeste Hoeve	52° 04'.95	5° 56'.94	6° 47'.0	18259y	18131y	2157y	43475y	+ 7y	+ 14y	- 16y
202	Scherpenzeel	52° 04'.66	5° 30'.60	7° 01'.4	18233y	18096y	2229y	43508y	- 10y	+ 7y	+ 21y
203	Bodegraven	52° 04'.59	4° 45'.95		18220y			43463y			- 23y
204	Stapelbroek	52° 04'.23	6° 22'.28					43519y			+ 36y
205	Ede	52° 04'.08	5° 44'.94	6° 59'.3	18262y	18126y	2222y	43492y	+ 4y	- 21y	+ 10y
206	Waddinxveen	52° 03'.82	4° 38'.20		18225y			43454y			- 26y
207	Beltrum	52° 03'.38	6° 34'.53		18287y			43487y			+ 11y
208	Oudenrijn	52° 03'.34	5° 03'.94		18251y			43432y			- 44y
209	Zegwaard	52° 03'.30	4° 30'.22					43444y			- 32y
210	Dieren	52° 03'.22	6° 07'.18	6° 40'.9	18262y	18138y	2125y	43457y	+ 43y	+ 20y	- 18y
211	Loosduinen	52° 02'.62	4° 14'.48	7° 41'.5	18202y	18038y	2436y	43444y	- 27y	- 9y	- 27y
212	Hengelo (Gld.)	52° 02'.43	6° 18'.05	6° 36'.3	18277y	18156y	2102y	43472y	- 4y	+ 16y	+ 4y
213	Remse	52° 01'.98	5° 23'.76					43537y			+ 73y
214	Doorn	52° 01'.76	5° 18'.42		18272y			43484y			+ 21y
215	Snelrewaard	52° 01'.64	4° 54'.79		18245y			43430y			- 32y
216	Werkhoven	52° 01'.60	5° 14'.87	7° 02'.2	18278y	18140y	2239y	43473y	+ 22y	+ 38y	+ 12y
217	Pijnacker	52° 01'.12	4° 25'.17	7° 37'.5	18235y	18074y	2420y	43423y	- 11y	- 19y	- 34y
218	Boven-Haastrecht	52° 00'.77	4° 49'.00					43423y			- 31y
219	Benschop	52° 00'.50	4° 58'.78		18265y			43411y			- 41y
220	Hummelo	52° 00'.44	6° 14'.26		18303y			43454y			+ 3y
221	Vreeswijk	51° 59'.62	5° 03'.67	7° 12'.3	18270y	18126y	2291y	43389y		+ 15y	- 56y
222	Zevenhuizen	51° 59'.50	4° 35'.02	7° 30'.2	18266y	18110y	2385y	43393y	+ 5y	- 8y	- 51y
223	Moordrecht	51° 59'.50	4° 40'.31					43410y			- 34y
224	Westervoort	51° 59'.25	6° 00'.40	6° 44'.9	18300y	18173y	2150y	43443y	+ 2y	+ 14y	+ 1y
225	Polsbroek	51° 59'.22	4° 54'.81		18264y			43407y			- 35y
226	Stolwijk	51° 58'.00	4° 45'.19		18270y			43394y			- 37y
227	Renkum	51° 57'.90	5° 42'.59	6° 53'.1	18309y	18178y	2195y	43425y	+ 10y	+ 15y	- 6y
228	Jaarsveld	51° 57'.84	4° 58'.28	7° 17'.2	18279y	18132y	2318y	43393y	- 4y	+ 2y	- 27y
229	Rijswijk	51° 57'.60	5° 21'.10	6° 57'.9	18276y	18141y	2216y	43472y	- 14y	+ 47y	+ 44y
230	Rhenen	51° 57'.40	5° 33'.25	7° 07'.0	18305y	18164y	2268y	43476y	- 1y	- 35y	+ 50y
231	Berkenwoude	51° 57'.30	4° 41'.52	7° 32'.1	18264y	18107y	2395y	43390y	- 20y	- 33y	- 36y
232	Gaag	51° 57'.25	4° 15'.46	7° 42'.8	18237y	18072y	2448y	43406y	- 37y	- 21y	- 19y
233	Zoelmond	51° 56'.87	5° 19'.09					43430y			+ 8y
234	Culemborg	51° 56'.79	5° 12'.78					43402y			- 19y
235	Woold	51° 56'.39	6° 42'.25	6° 33'.5	18353y	18233y	2096y	43426y	- 8y	- 32y	+ 8y
236	Aalten	51° 55'.97	6° 34'.68	6° 30'.9	18347y	18228y	2082y	43419y	+ 5y	- 2y	+ 5y
237	Schoonhoven	51° 55'.81	4° 49'.24	7° 30'.5	18272y	18166y	2388y	43377y	- 29y	- 44y	- 37y
238	Hettenheuvel	51° 55'.28	6° 14'.10	6° 40'.2	18327y	18203y	2129y	43413y	- 10y	+ 3y	+ 4y
239	Schoonrewoerd	51° 55'.26	5° 06'.74					43389y			- 20y
240	Ouderkerk	51° 55'.22	4° 39'.52		18264y			43359y			- 50y
241	Oostvoorne	51° 54'.91	4° 07'.15	7° 47'.5	18241y	18072y	2473y	43374y	- 49y	- 23y	- 32y
242	Silvolde	51° 54'.35	6° 22'.92	6° 37'.6	18350y	18228y	2118y	43413y	+ 2y	- 7y	+ 12y
243	Noordeloos	51° 53'.97	4° 57'.37	7° 21'.2	18306y	18156y	2343y	43366y	- 9y	- 18y	- 32y
244	Aerd	51° 53'.70	6° 05'.97	6° 45'.9	18331y	18203y	2159y	43398y	- 15y	- 6y	+ 2y
245	Druten	51° 53'.60	5° 36'.20					43432y			+ 37y
246	Winsum	51° 53'.57	5° 41'.28	7° 05'.9	18334y	18194y	2266y	43407y	- 7y	- 51y	+ 13y
247	Ottoland	51° 53'.54	4° 52'.47					43346y			- 48y
248	Beesd I	51° 53'.24	5° 12'.10	7° 05'.4				43337y			- 55y
249	Gent	51° 52'.60	5° 56'.29	6° 47'.9	18360y	18230y	2173y	43382y	+ 10y	+ 5y	- 4y
250	Wadenhoijen	51° 52'.52	5° 21'.28					43340y			- 46y
251	Zwartewaal	51° 52'.50	4° 13'.70	7° 42'.7	18240y	18076y	2448y	43364y	- 68y	- 13y	- 22y
252	Alblasserdam	51° 52'.39	4° 40'.48	7° 33'.9	18291y	18131y	2408y	43363y	- 34y	- 40y	- 22y
253	Tiel	51° 52'.32	5° 25'.12	6° 57'.9	18361y	18225y	2226y	43363y	+ 26y	+ 30y	- 21y

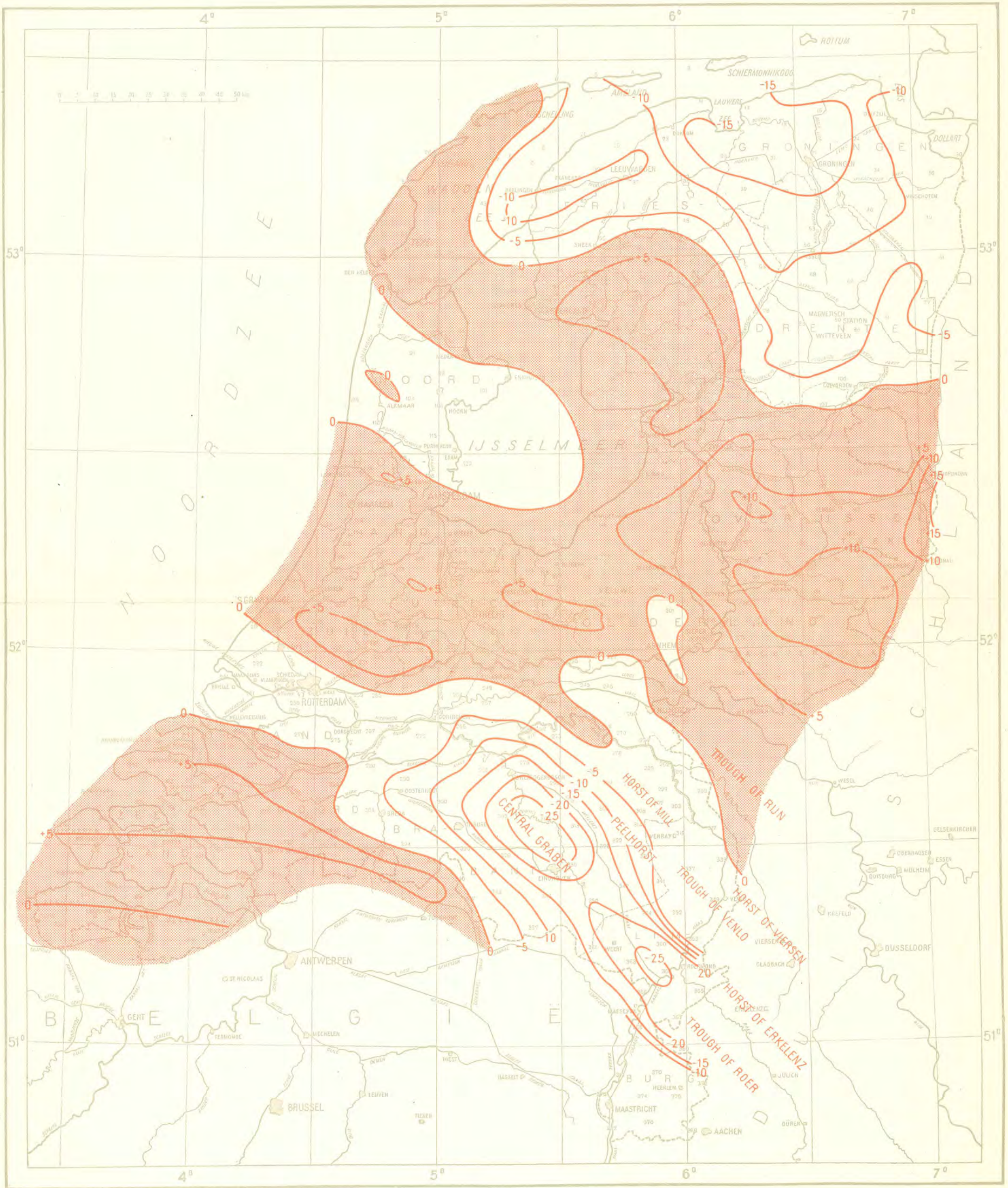
	Name of station	φ	λ	-D	H	X	-Y	Z	ΔX	ΔY	ΔZ
339	Rilland-Bath	51° 24'.58	4° 12'.48	7° 44'.1	18419 γ	18252 γ	2479 γ	43279 γ	-110 γ	- 27 γ	+123 γ
340	De Utrecht	51° 24'.31	5° 05'.22	7° 24'.9	18501 γ	18346 γ	2388 γ	43133 γ	- 58 γ	- 68 γ	- 18 γ
341	Helenaveen	51° 23'.52	5° 54'.23	6° 33'.7	18537 γ	18416 γ	2118 γ	43168 γ	- 31 γ	+ 79 γ	+ 24 γ
342	Heusden	51° 22'.97	5° 45'.38	6° 43'.3	18533 γ	18406 γ	2169 γ	43126 γ	- 38 γ	+ 52 γ	- 13 γ
343	Bladel	51° 22'.31	5° 14'.27	7° 17'.5	18530 γ	18380 γ	2352 γ	43089 γ	- 47 γ	- 54 γ	- 45 γ
344	Valkenswaard	51° 21'.68	5° 26'.73	7° 03'.7	18455 γ	18405 γ	2280 γ	43074 γ	- 35 γ	- 13 γ	- 53 γ
345	Cadzand	51° 21'.62	3° 24'.05	7° 59'.4	18482 γ	18303 γ	2569 γ	43286 γ	- 46 γ	+ 5 γ	+158 γ
346	Hengstdijk	51° 21'.62	4° 01'.18	7° 57'.0				43255 γ			+127 γ
347	Reusel	51° 21'.22	5° 07'.88	7° 23'.4	18528 γ	18374 γ	2383 γ	43096 γ	- 56 γ	- 69 γ	- 29 γ
348	Schoondijke	51° 21'.18	3° 34'.79	7° 59'.1	18463 γ	18284 γ	2565 γ	43309 γ	- 76 γ	- 18 γ	+185 γ
349	Tegelen	51° 20'.77	6° 07'.28	6° 44'.2	18547 γ	18419 γ	2176 γ	43134 γ	- 58 γ	+ 9 γ	+ 13 γ
350	Sterksel	51° 20'.70	5° 37'.65	6° 51'.5	18546 γ	18413 γ	2215 γ	43076 γ	- 43 γ	+ 26 γ	- 44 γ
351	Nieuw-Namen	51° 18'.96	4° 10'.18	7° 46'.6	18486 γ	18316 γ	2501 γ	43230 γ	- 88 γ	- 40 γ	+124 γ
352	Helden	15° 18'.84	5° 58'.24	6° 40'.0	18553 γ	18427 γ	2154 γ	43144 γ	- 59 γ	+ 36 γ	+ 41 γ
353	Zaamslag	51° 18'.70	3° 52'.38	7° 50'.4	18516 γ	18342 γ	2526 γ	43249 γ	- 51 γ	- 21 γ	+145 γ
354	Nederweert	51° 18'.65	5° 50'.83	6° 43'.3	18565 γ	18438 γ	2173 γ	43110 γ	- 44 γ	+ 36 γ	+ 6 γ
355	Biervliet	51° 17'.92	3° 41'.39	7° 53'.7	18511 γ	18375 γ	2548 γ	43285 γ	- 16 γ	- 15 γ	+188 γ
356	Aardenburg	51° 16'.62	3° 28'.80	8° 02'.7	18567 γ	18384 γ	2598 γ	43203 γ	- 8 γ	- 33 γ	+116 γ
357	Schaft	51° 16'.36	5° 23'.51	7° 03'.7	18572 γ	18431 γ	2283 γ	43033 γ	- 49 γ	- 5 γ	- 52 γ
358	Hulst	51° 15'.06	4° 00'.20	7° 45'.8	18531 γ	18361 γ	2503 γ	43274 γ	- 66 γ	- 16 γ	+200 γ
359	Swalmen	51° 15'.05	6° 02'.52	6° 38'.5	18560 γ	18436 γ	2147 γ	43160 γ	- 83 γ	+ 33 γ	+ 86 γ
360	Heythuizen	51° 14'.49	5° 54'.23	6° 36'.4	18486 γ	18462 γ	2138 γ	43086 γ	- 56 γ	+ 64 γ	+ 17 γ
361	Dorplein	51° 13'.88	5° 37'.57	6° 54'.1	18575 γ	18440 γ	2232 γ	43022 γ	- 69 γ	+ 12 γ	- 42 γ
362	Sas van Gent	51° 13'.52	3° 49'.76	7° 40'.8	18594 γ	18427 γ	2485 γ	43193 γ	- 4 γ	+ 29 γ	+132 γ
363	Hunsel	51° 11'.73	5° 47'.13	6° 45'.5	18579 γ	18450 γ	2186 γ	43019 γ	- 83 γ	+ 35 γ	- 26 γ
364	Luinc	51° 10'.00	5° 55'.39	6° 35'.3	18533 γ	18510 γ	2138 γ	43043 γ	- 44 γ	+ 64 γ	+ 12 γ
365	Posterholt	51° 06'.96	6° 03'.40	6° 24'.8	18539 γ	18522 γ	2082 γ	43141 γ	- 62 γ	+101 γ	+135 γ
366	Ohé	51° 06'.86	5° 49'.16	6° 47'.2	18598 γ	18468 γ	2198 γ	43027 γ	-106 γ	+ 21 γ	+ 22 γ
367	Koningsbosch	51° 03'.04	5° 57'.72	6° 37'.4	18687 γ	18562 γ	2155 γ	43034 γ	- 47 γ	+ 44 γ	+ 60 γ
368	Born	51° 01'.70	5° 47'.82	6° 48'.9	18668 γ	18536 γ	2215 γ	43006 γ	- 77 γ	+ 9 γ	+ 44 γ
369	Schinveld	50° 58'.60	5° 59'.62	6° 41'.8	18762 γ	18634 γ	2188 γ	42988 γ	- 12 γ	+ 9 γ	+ 51 γ
370	Nuth	50° 54'.54	5° 53'.03	6° 53'.2	18795 γ	18660 γ	2254 γ	42977 γ	- 13 γ	- 39 γ	+ 74 γ
371	Geulle	50° 54'.51	5° 43'.95	7° 05'.0	18740 γ	18597 γ	2311 γ	43061 γ	- 69 γ	- 73 γ	+158 γ
372	Kerkrade	50° 52'.82	6° 03'.67	6° 45'.8	18798 γ	18667 γ	2214 γ	42908 γ	- 27 γ	- 24 γ	+ 19 γ
373	Maastricht	50° 51'.52	5° 38'.63	7° 19'.9	18822 γ	18668 γ	2402 γ	43106 γ	- 18 γ	-150 γ	+229 γ
374	Valkenburg	50° 51'.03	5° 49'.23	6° 47'.9	18834 γ	18701 γ	2229 γ	42909 γ	+ 3 γ	- 4 γ	+ 38 γ
375	Ubachsberg	50° 50'.78	5° 57'.21	6° 48'.6	18805 γ	18673 γ	2230 γ	42892 γ	- 32 γ	- 23 γ	+ 20 γ
376	Slenaken	50° 46'.96	5° 50'.62	6° 59'.2	18790 γ	18651 γ	2286 γ	42845 γ	- 79 γ	- 62 γ	+ 6 γ
377	Eijsden	50° 45'.87	5° 41'.54	7° 09'.2	18826 γ	18680 γ	2344 γ	42817 γ	- 52 γ	- 96 γ	- 14 γ
378	Vaals	50° 45'.84	6° 00'.74	6° 39'.9	18801 γ	18674 γ	2182 γ	42887 γ	- 73 γ	+ 18 γ	+ 58 γ
A	Oosterom	53° 20'.70	5° 16'.20	(7° 18')	17651 γ	17508 γ	2243 γ	44211 γ	+ 8 γ	- 9 γ	+ 90 γ
B	N. O. Meep	53° 18'.50	5° 26'.30	(7° 18')	17722 γ	17578 γ	2252 γ	44149 γ	+ 53 γ	- 42 γ	+ 47 γ
C	Franekegat	53° 17'.40	5° 16'.30	(7° 25')	17701 γ	17552 γ	2285 γ	44181 γ	+ 26 γ	- 50 γ	+ 88 γ
D	Schieringhals	53° 16'.80	5° 09'.70	(7° 18')	17662 γ	17519 γ	2244 γ	44213 γ	- 6 γ	+ 9 γ	+124 γ
E	N. O. Ballastplaat	53° 15'.75	5° 26'.20	(7° 17')	17755 γ	17612 γ	2251 γ	44104 γ	+ 66 γ	- 40 γ	+ 24 γ
F	Oost v. Griend	53° 14'.50	5° 17'.50	(7° 26')	17743 γ	17594 γ	2295 γ	44136 γ	+ 44 γ	- 61 γ	+ 67 γ
G	Z. Ballastplaat	53° 13'.40	5° 23'.10	(7° 17')	17757 γ	17614 γ	2254 γ	44083 γ	+ 52 γ	- 34 γ	+ 23 γ
H	W. v. Inschot	53° 13'.00	5° 08'.00	(7° 26')	17717 γ	17568 γ	2292 γ	44261 γ	+ 14 γ	- 33 γ	+205 γ
I	Lange Zand	53° 11'.50	5° 12'.10					44247 γ			+203 γ
J	N. v. Scheurak	53° 06'.30	5° 03'.20	(7° 19')	17882 γ	17736 γ	2280 γ	44071 γ	+133 γ	- 6 γ	+ 71 γ
K	Einde Doove Balg	53° 04'.90	5° 16'.25					43993 γ			+ 4 γ
L	Stompe Waard	53° 02'.90	5° 07'.10					43984 γ			+ 12 γ
M	Harde Bollen	53° 00'.33	4° 51'.65	(7° 17')	17863 γ	17719 γ	2265 γ	43926 γ	+ 78 γ	+ 41 γ	- 25 γ
N	Lutjeswaard	52° 59'.70	5° 02'.70	(7° 17')	17881 γ	17737 γ	2267 γ	43921 γ	+ 82 γ	+ 11 γ	+ 12 γ



Δ -CHART. GEOMAGNETIC ANOMALIES IN THE NETHERLANDS. RED ARE POSITIVE REGIONS (TOO LARGE VALUE OF THE VERTICAL COMPONENT), WHITE ARE NEGATIVE REGIONS (TOO SMALL VALUE OF THE VERTICAL COMPONENT). THE FIGURES GIVE VALUES OF ΔZ (IN γ), THE VECTORS INDICATE THE DIRECTION AND THE MAGNITUDE OF THE HORIZONTAL DEVIATIONS. THE ISANOMALS OF ΔZ HAVE BEEN DRAWN FOR EVERY 20 γ .

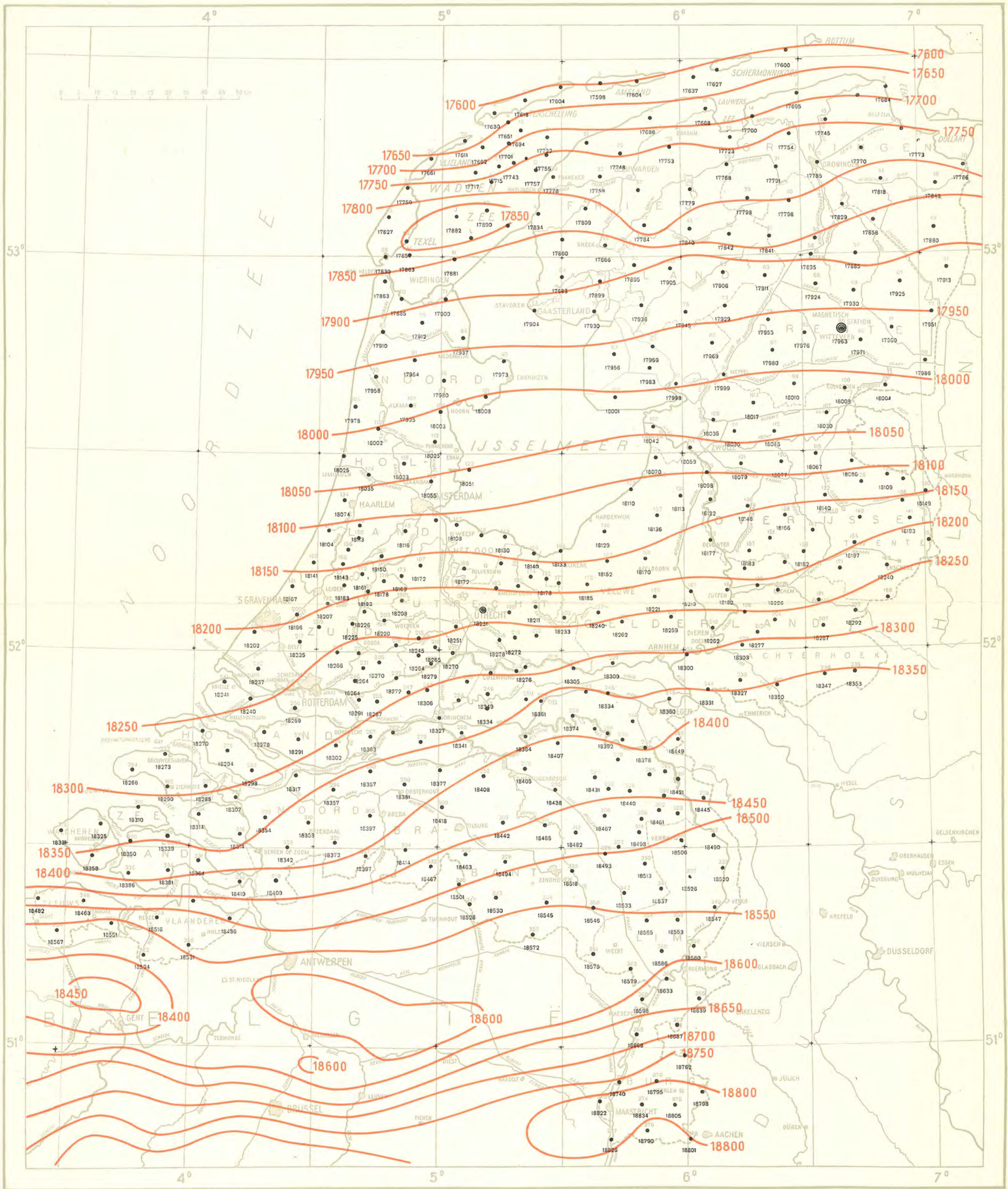


D-CHART. WESTERLY DECLINATION OF THE GEOMAGNETIC FIELD IN THE NETHERLANDS, EPOCH 1945.0. ANNUAL DECREASE ABOUT 8'.

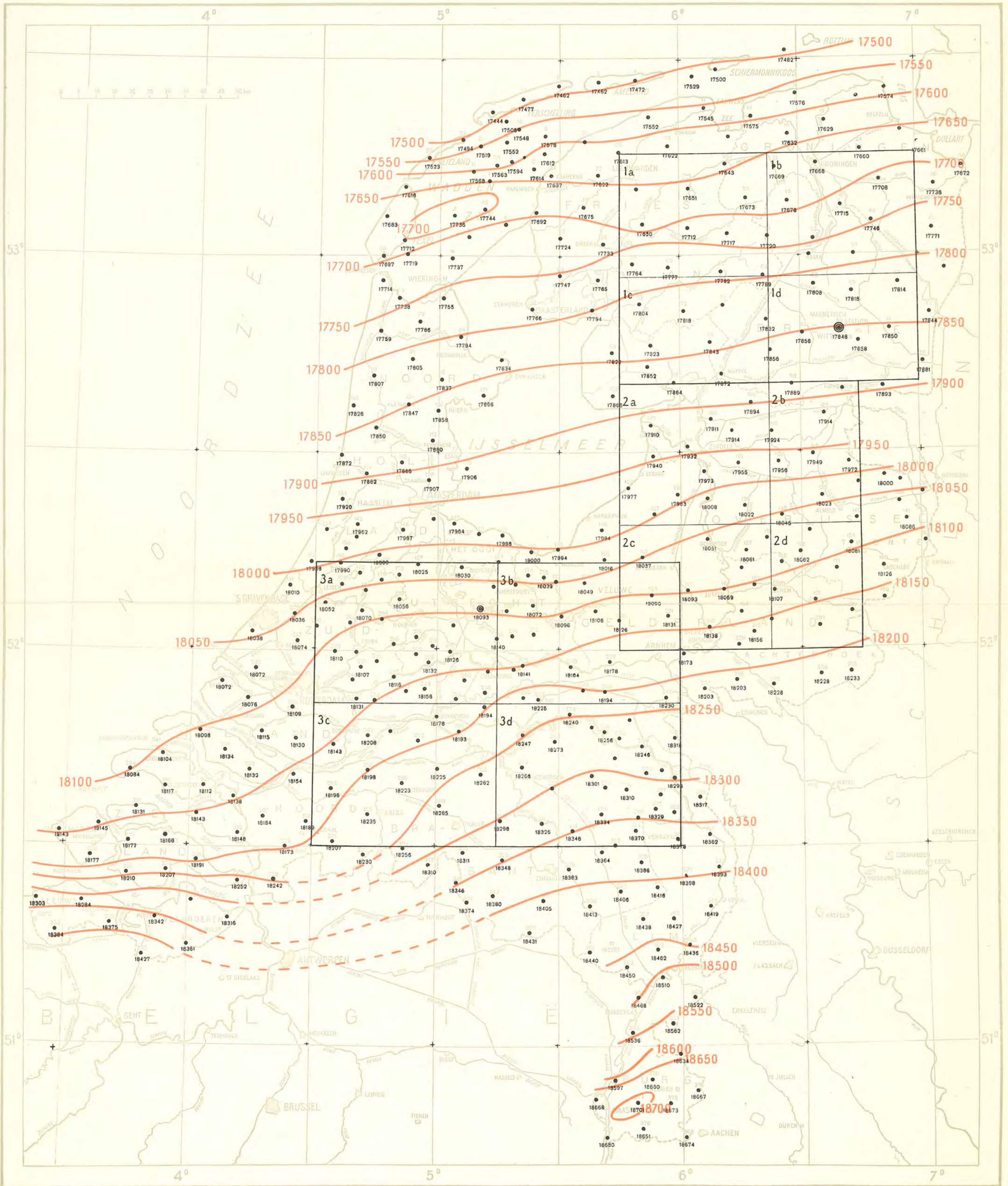


G-CHART. GRAVITY ANOMALIES (IN MGAL) IN THE NETHERLANDS. RED ARE POSITIVE REGIONS, WHITE ARE NEGATIVE REGIONS. THE ISOGAMS HAVE BEEN DRAWN FOR EVERY 5 MGAL.

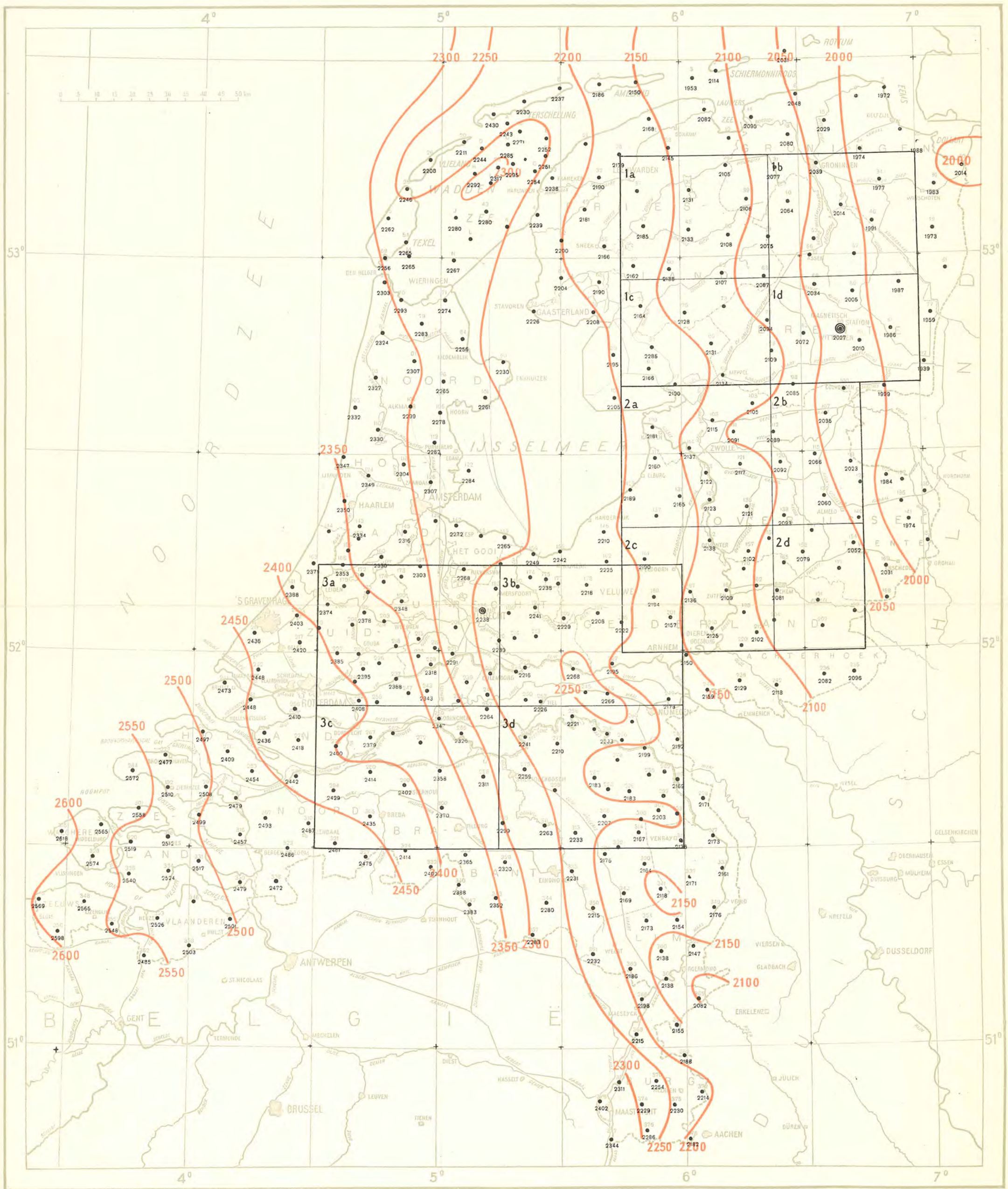
This chart was supplied by the Bataafse Petroleum Maatschappij.



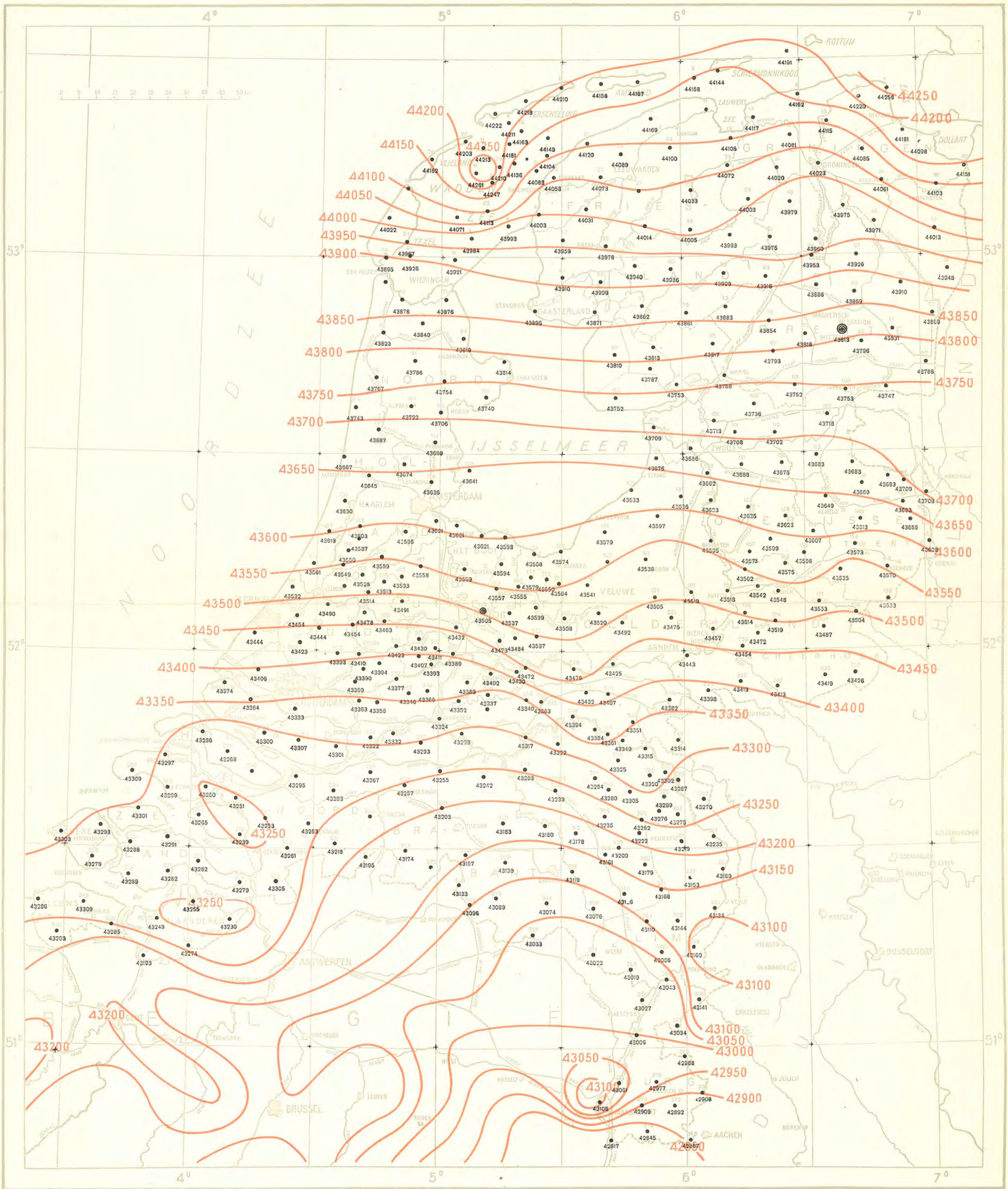
H-CHART. HORIZONTAL INTENSITY (IN γ) OF THE GEOMAGNETIC FIELD IN THE NETHERLANDS, EPOCH 1945.0. ANNUAL CHANGE ABOUT 0 γ .



X-CHART. NORTHERLY COMPONENT (IN γ) OF THE GEOMAGNETIC FIELD IN THE NETHERLANDS, EPOCH 1945.0. ANNUAL CHANGE ABOUT 0 γ .
 IN THE TRAPEZIA THE AIR-EARTH CURRENTS HAVE BEEN COMPUTED.



Y-CHART. WESTERLY COMPONENT (IN γ) OF THE GEOMAGNETIC FIELD IN THE NETHERLANDS, EPOCH 1945.0. ANNUAL DECREASE ABOUT 50 γ . IN THE TRAPEZIA THE AIR-EARTH CURRENTS HAVE BEEN COMPUTED.



Z-CHART. VERTICAL COMPONENT (IN γ) OF THE GEOMAGNETIC FIELD IN THE NETHERLANDS, EPOCH 1945.0. ANNUAL INCREASE ABOUT 25 γ .

